



Community Workshop on Air Quality Remote Sensing from Space: Defining an Optimum Observing Strategy

Directly Retrieved Global Distribution of Tropospheric Column Ozone from GOME

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Abstract Ozone profiles and tropospheric column ozone are derived from back scattered radiance spectra in the ultraviolet measured by the GOME with the optimal estimation technique. The retrievals are capable of capturing the spatiotemporal evolution of TCO in response to regional or short time-scale events such as the 1997-1998 El Niño event and a 10-20 DU change within a few days. The mean biases relative to ozonesonde observations are usually within 3 DU (15%) and the standard deviations are within 3-8 DU (13-27%). The global distribution of TCO displays nearly zonal bands of enhanced TCO of 36-48 DU at 20°S-30°S during the austral spring and at 25°N-45°N during the boreal spring and summer. The overall structure is very similar to the GEOS-CHEM model simulation with small biases of <±5 DU and consistent seasonal cycles in most regions. This method can be readily applied to observations from current or next generation of nadir-viewing UV/visible instruments such as SCIAMACHY, OMI, GOME-2, and OMPS to derive ozone profiles, total ozone and tropospheric ozone.

Station	TCO			Station	TCO								
	A	Priori-Sonde	GOME-Sonde		A	Priori-Sonde	GOME-Sonde						
Resolute	5.6	3.7	0.56	0.1	6.2	0.46	Ku. Lumpur	9.0	5.0	0.00	3.0	5.4	0.38
Scoreby	2.4	5.4	0.52	-3.3	5.9	0.39	San Crist.	7.8	2.7	0.74	2.7	3.8	0.62
Sodankylä	2.7	3.8	0.62	-0.5	6.5	0.53	Nairobi	-1.7	5.9	0.29	-0.7	5.4	0.66
Churchill	3.1	5.2	0.56	1.0	4.2	0.62	Java	-4.6	7.5	0.50	-0.1	5.0	0.87
Valencia	-0.3	8.0	0.27	0.5	6.5	0.25	Ascension	-6.9	6.3	0.80	0.0	5.9	0.77
Hohenpei	2.4	5.3	0.59	1.1	5.8	0.56	Am. Samoa	5.0	5.8	0.55	1.5	4.5	0.82
Payenne	2.3	5.6	0.64	-0.1	5.6	0.65	Tahiti	9.6	5.6	0.47	5.1	4.1	0.82
Boulder	3.1	5.7	0.69	1.9	3.9	0.87	Reunion	1.6	5.3	0.78	1.8	3.2	0.85
Ankara	3.4	4.9	0.83	2.6	5.2	0.87	Irene	3.6	5.5	0.79	-0.3	4.6	0.85
Wallop Is.	0.6	6.6	0.69	-3.0	5.4	0.83	Easter Is.	7.2	8.3	0.45	2.6	5.4	0.81
Tateno	3.5	6.6	0.73	0.0	5.2	0.85	Laverton	0.8	3.4	0.68	2.5	4.7	0.69
Kaoshima	6.2	7.6	0.65	1.7	7.8	0.71	Lauder	2.3	3.5	0.62	1.6	4.0	0.40
Santa Cruz	-4.2	8.2	0.54	1.5	7.4	0.76	Macquarie	2.6	3.2	0.75	-0.3	5.2	0.56
Naha	6.4	7.0	0.56	5.4	6.1	0.70	Marambio	-0.7	5.1	0.76	-0.9	5.0	0.81
Hilo	2.6	6.4	0.67	1.7	6.0	0.79	Syowa	1.0	3.8	0.92	-3.1	5.6	0.80
Panamaribo	2.0	3.7	0.00	-1.8	3.7	0.22	Neumayer	1.7	3.4	0.93	-4.0	5.1	0.89
Kaashidhoo	3.4	5.0	0.27	0.9	6.5	0.27							

Table 1. Comparison statistics (mean bias in DU), 1σ standard deviation in DU, and the Pearson correlation coefficient) between GOME retrieved/a priori TCO and ozonesonde measurements at 33 stations from 1996 though 1999.

• Mean biases are mostly within 3 DU (15%) and the 1σ standard deviations are within 3-8 DU 13-27%.

1. Algorithm Description

- Optimal estimation with TOMS V8 ozone climatology
- Fitting window: 289-307 and 326-339 nm (couldn't be used in between due to calibration problems).
- 11 layers (2-3 layers in the troposphere, use NCEP tropopause to divide troposphere and stratosphere)
- Wavelength calibrations: variable slit widths, shifts among radiances, irradiances, and cross-sections
- Radiometric calibrations: undersampling correction, a 2nd-order polynomial correction in 290-310 nm and a 2nd-order surface albedo in 325-339 nm,

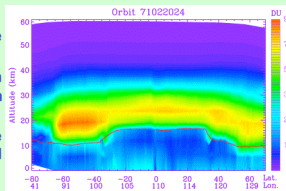


Fig. 1 An orbit of retrieved profiles of partial (layer) column ozone (DU) on 22 October 1997. It shows ozone hole in Antarctic polar vortex and maximum ozone outside the polar vortex, enhanced tropospheric ozone over Indonesia and the Southern India Ocean.

Fig. 1 was made with JROBT, FOMF, NCEP reanalysis, temperature, surface & tropopause pressure, GOMECAL, clouds, GOCART

2. Retrieval Characterization of Tropospheric Ozone

- Vertical resolution: 9-16 km in the troposphere DFS: 0.5-1.5 between ±50°N/S.
- Precision: < 10% Precision + Smoothing: 15-25%
- Precision in TCO: from <6% (1.5 DU) in the tropics to <12% (3.0 DU) at high latitudes
- Precision + smoothing in TCO: from <12% in tropics to <25% at high latitudes.
- Globally-averaged retrieval accuracy in TCO: 21%.
- A priori influence in TCO: from ~15% in the tropics to ~50% at high latitudes.
- Spatial resolution: normally 960x80 km² (320x40 km² is possible after June 1998)

4. Global Distribution of TCO

- Tropical wave-1 during most months especially in the southern hemisphere.
- TCO of 33-39 DU near Gulf of Guinea during North Africa biomass burning season (support the tropical Atlantic paradox).
- Significant spatiotemporal variation over tropical pacific, location of low TCO migrates with ITCZ.
- Wave-1 patterns weakens in the extratropics with nearly uniform bands at ~30°S/N.
- Enhanced TCO at 25°N-45°N in April-July (36-48 DU) and at 25°S-35°S in Sep.-Nov. (33-45 DU).
- At ~30°S and over the tropical south Atlantic, TCO maximizes in Sep.-Nov. and minimizes in Apr.-June.
- 15-30°N: highest in spring and lowest in Jul.-Sep. 30-45°N: highest in May-Jul. and lowest in Nov.-Dec.
- TCO of <30 DU uniformly distributed south of 35°S during all seasons with a large latitudinal

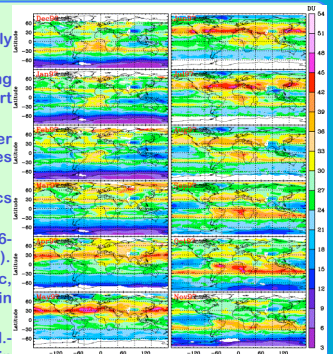


Fig. 5 Global maps of monthly mean TCO from Dec. 1997 to Nov. 1997.

5. Comparison with the GEOS-CHEM

- Similar overall structure: tropical wave-1 pattern, more uniform distribution of enhanced TCO at ~30°N/S, generally similar spatiotemporal variations at middle to high latitudes.
- Seasonal average difference:
 - Global: <±2.4 DU, r: 0.8-0.9
 - SH: 1.2±2.1 DU, r: 0.94-0.98
 - NH: 4.3±4.6 DU, r: 0.6-0.8
- Significant discrepancies remain:
 - In GEOS-CHEM, the northern bands are broader and are extended to some subtropical and tropical areas; GOME TCO is consistently smaller by 5-20 DU over these regions.
 - GOME TCO does not show a regional maximum of >42 DU over India and Southeast Asia in MAM and Middle-East maximum in JJA and SON, leading to 5-20 DU differences.

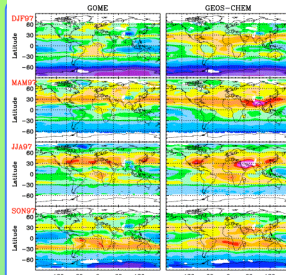


Fig. 6 Comparison of GOME and GEOS-CHEM TCO for four seasons during Dec 1996-Nov 1997 (DJF, MAM, JJA, SON).

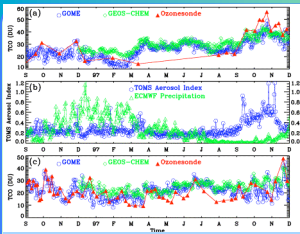


Fig. 2 (a) GOME and GEOS-CHEM daily TCO from Sep. 1996 to Nov. 1997 over Indonesia (110°-125°E, 8.6°-6.6°S) and ozonesonde TCO at Java (112.7°E, 7.6°S). (b) Same as (a) but for TOMS aerosol index and ECMWF precipitation. (c) Same as (a) but over (180°-158°W, 15.2°-13.2°S) and at American Samoa (170°W, 14.2°S).

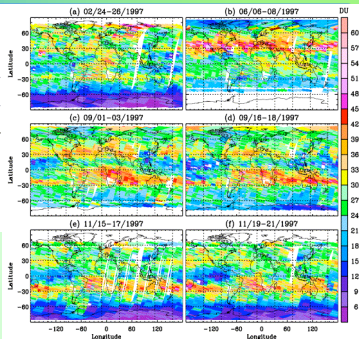


Fig. 3 Three-day composite global maps of TCO.

3. Daily Retrievals and Validation

- Detailed TCO response to the 1997-1998 El Niño event (dry weather and wildfires) over Indonesia (Fig. 2a-b).
- Changes of 10-20 DU within a few days (Fig. 2c).
- TCO (>40 DU) in Feb. over North Africa (Fig. 3a)
- Band of high TCO (40-60 DU) at 25-45°N in June (Fig. 3b).
- Distinct differences at 25-40°N (Fig. 3c-d) between 1-3 Sep. and 16-18 Sep.; a 12 DU increase over Indonesia.
- Transport of southern subtropical high-ozone air to the tropics around American Samoa in Nov. (Fig. 2c, Fig. 3e-f).
- Capture most of the temporal variability of ozonesonde observations (Fig. 2a, Fig. 2c, Fig. 3).

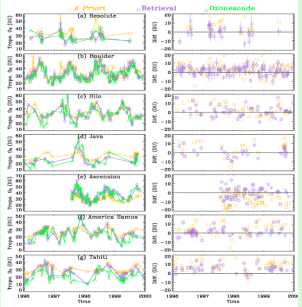


Fig. 4 Comparison with selected ozonesonde observations.

Conclusions and Future Outlook

- Ozone profiles and tropospheric column ozone are derived from GOME data using the optimal estimation technique after detailed wavelength and radiometric calibrations and forward modeling.
- The retrieved TCO captures most of the temporal variability in daily ozonesonde observations with mean biases usually within 3 DU (15%).
- Both GOME and GEOS-CHEM TCO show similar overall structures: tropical wave-1 structure, decline of the wave-1 structure in the extratropics, nearly zonal bands of enhanced TCO of 36-45 DU at 20°S-30°S during the austral spring and at 25°N-45°N during boreal spring and summer, TCO of <30 DU zonally distributed at southern middle to high-latitudes.
- Significant positive biases of 5-20 DU occur at some northern tropical and subtropical regions.
- Reprocessing of GOME data (July 1995-June 2003) is ongoing with a degradation correction scheme, improvements in trace gas (SO₂, BrO, NO₂, HCHO) fitting (e.g., fitting weighting functions, using model-simulated profiles to represent a priori and derive effect air mass factor), and other minor enhancements.
- Apply this algorithm to SCIAMACHY, OMI and GOME-2 observations.

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See the following papers for more detail:
Liu, X., K. Chance, C.E. Sioris, R.J.D. Spurr, T.P. Kurosu, R.V. Martin, M.J. Newchurch, Ozone Profile and Tropospheric Ozone Retrieval from Global Ozone Monitoring Experiment (GOME): Algorithm Description and Validation, J. Geophys. Res., 110(D20), D20307, 10.1029/2005JD006240, 2005.
Liu et al., First directly-retrieved global distribution of tropospheric column ozone: comparison with the GEOS-CHEM model, J. Geophys. Res., Vol. 111, No. D2, D02308, 10.1029/2005JD006564, 2006.