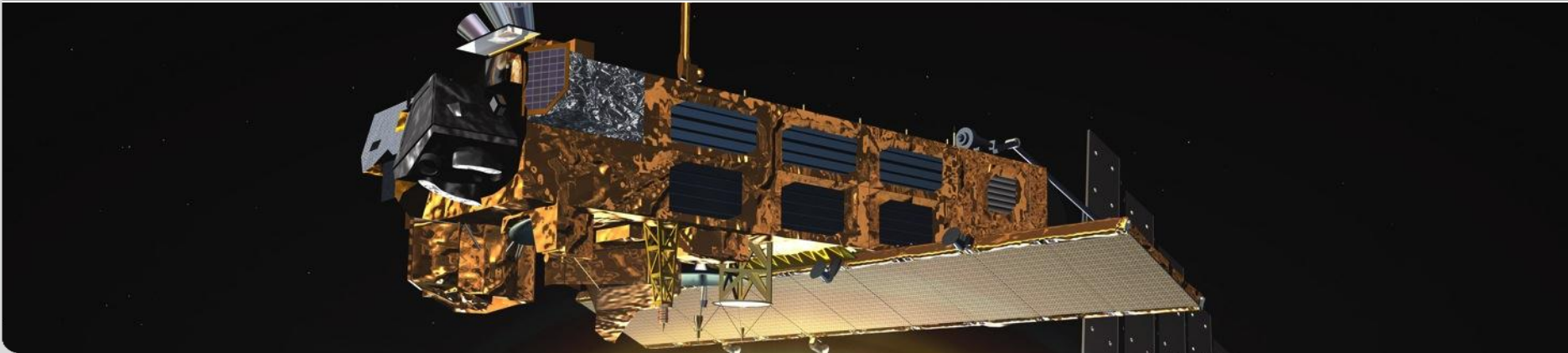


MIPAS observations of the Asian monsoon anticyclone: Focus on water vapor, tracers, and pollutants

Gabriele Stiller, Norbert Glatthor, Michael Höpfner, Stefan Lossow, Andreas Wiegele,
and Thomas von Clarmann

Institute of Meteorology and Climate Research (IMK-ASF)



MIPAS overview

- FTIR limb
2002 to A
- Altitude c
- More than
temperatu
- Full data i
see <http://>
- Trace spe
CO, HCN
- Several re
100 000 C

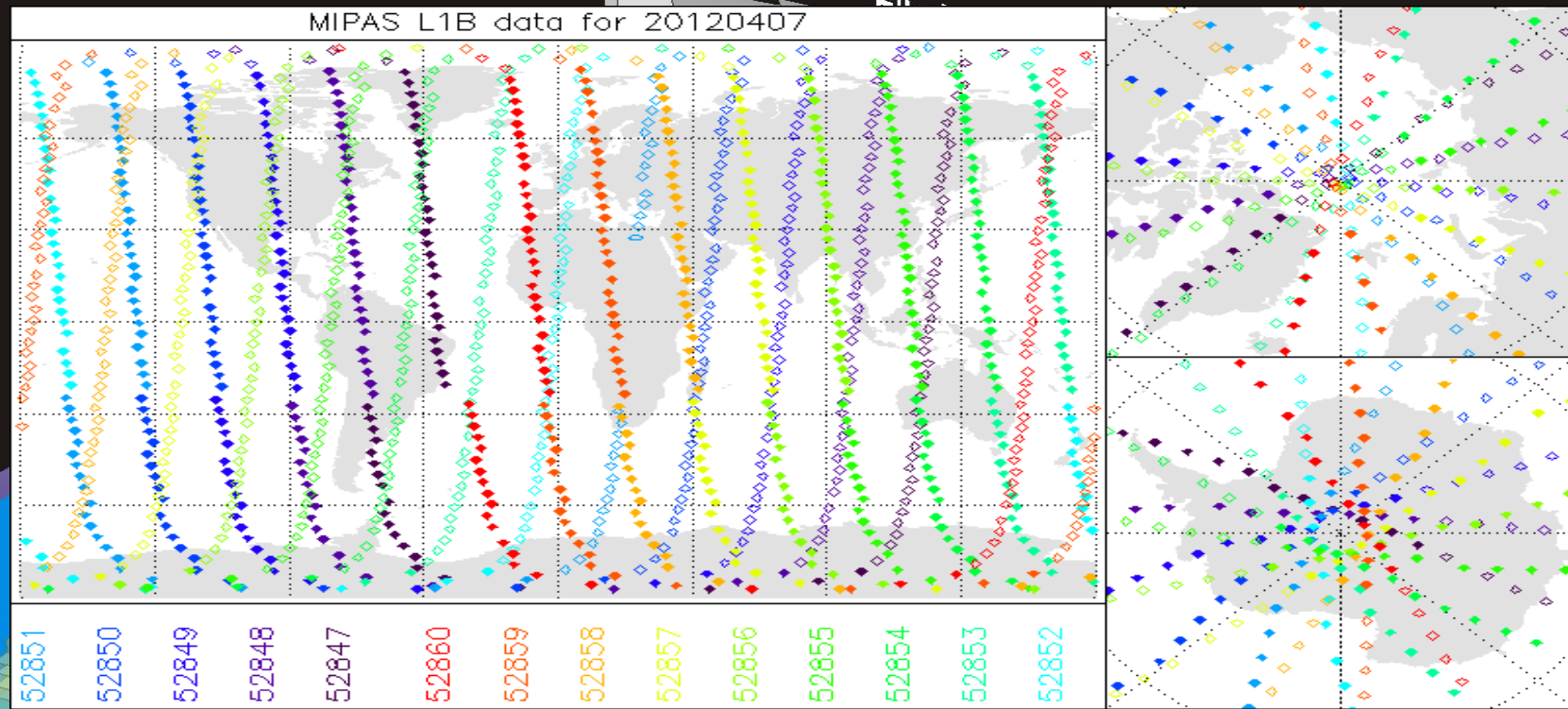


visat from July

2 km
clouds and
000 before 2005)
ersion 5) available,
IPAS data IMK“
O₃, C₂H₂, C₂H₆,

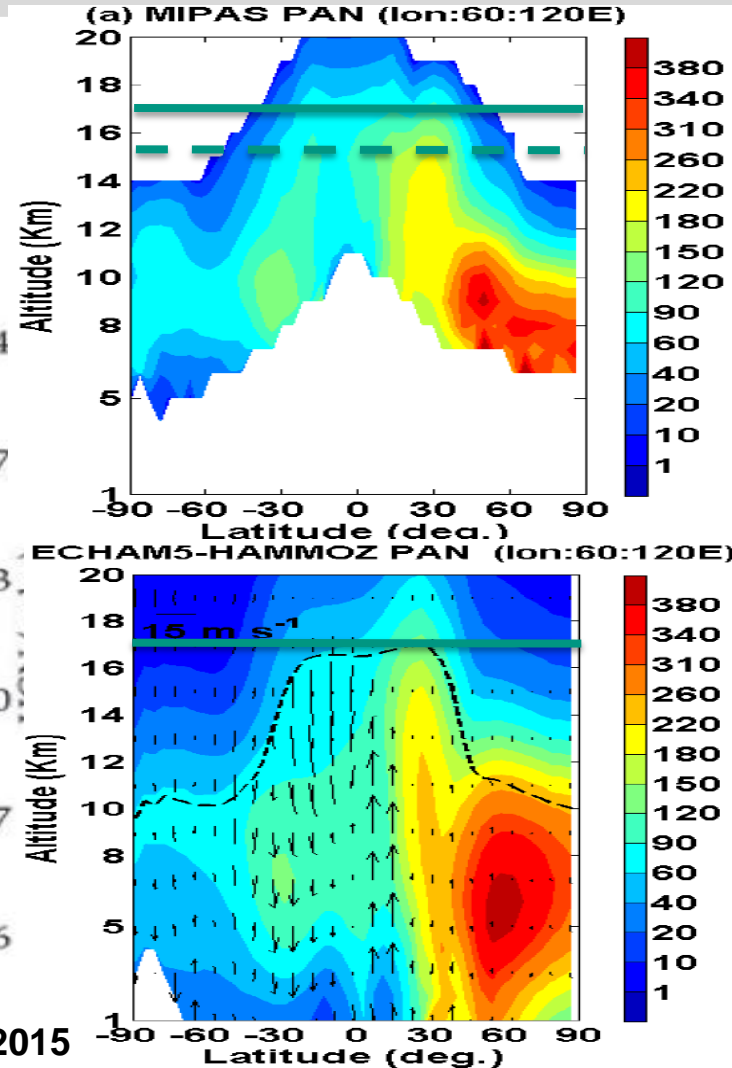
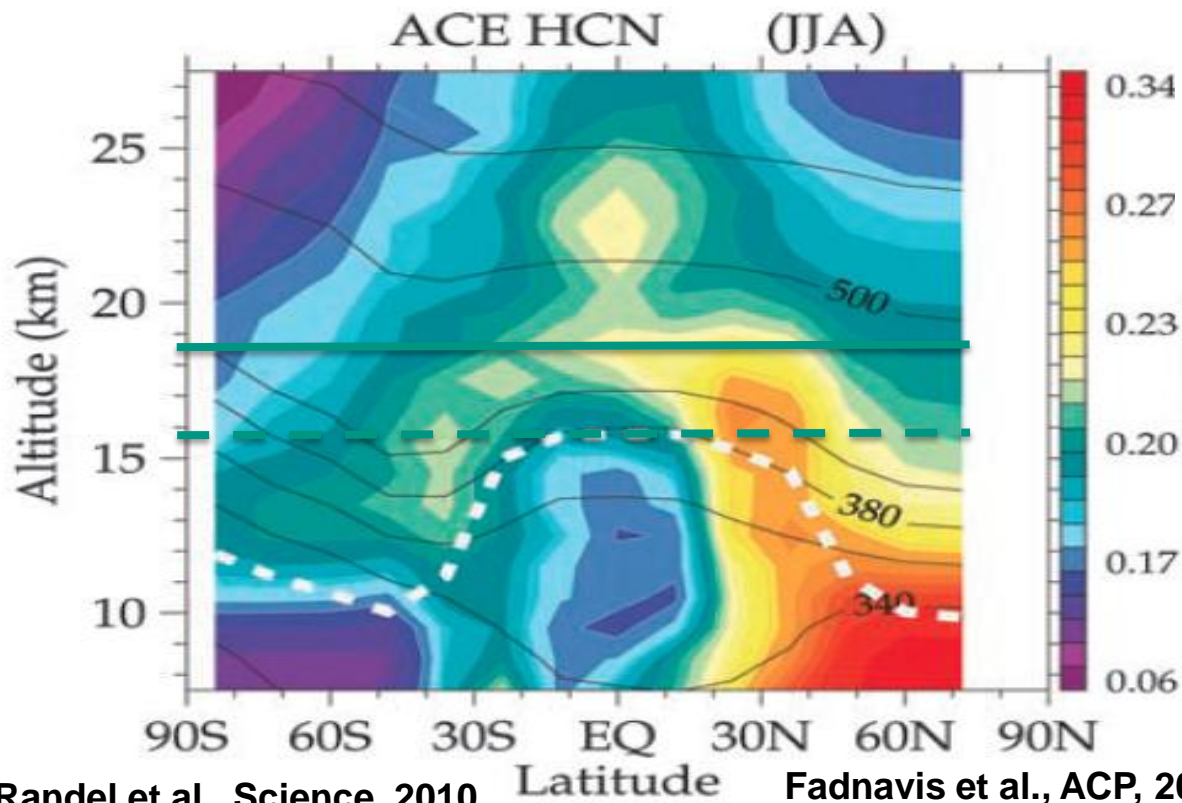
e species → the

MIPAS overview

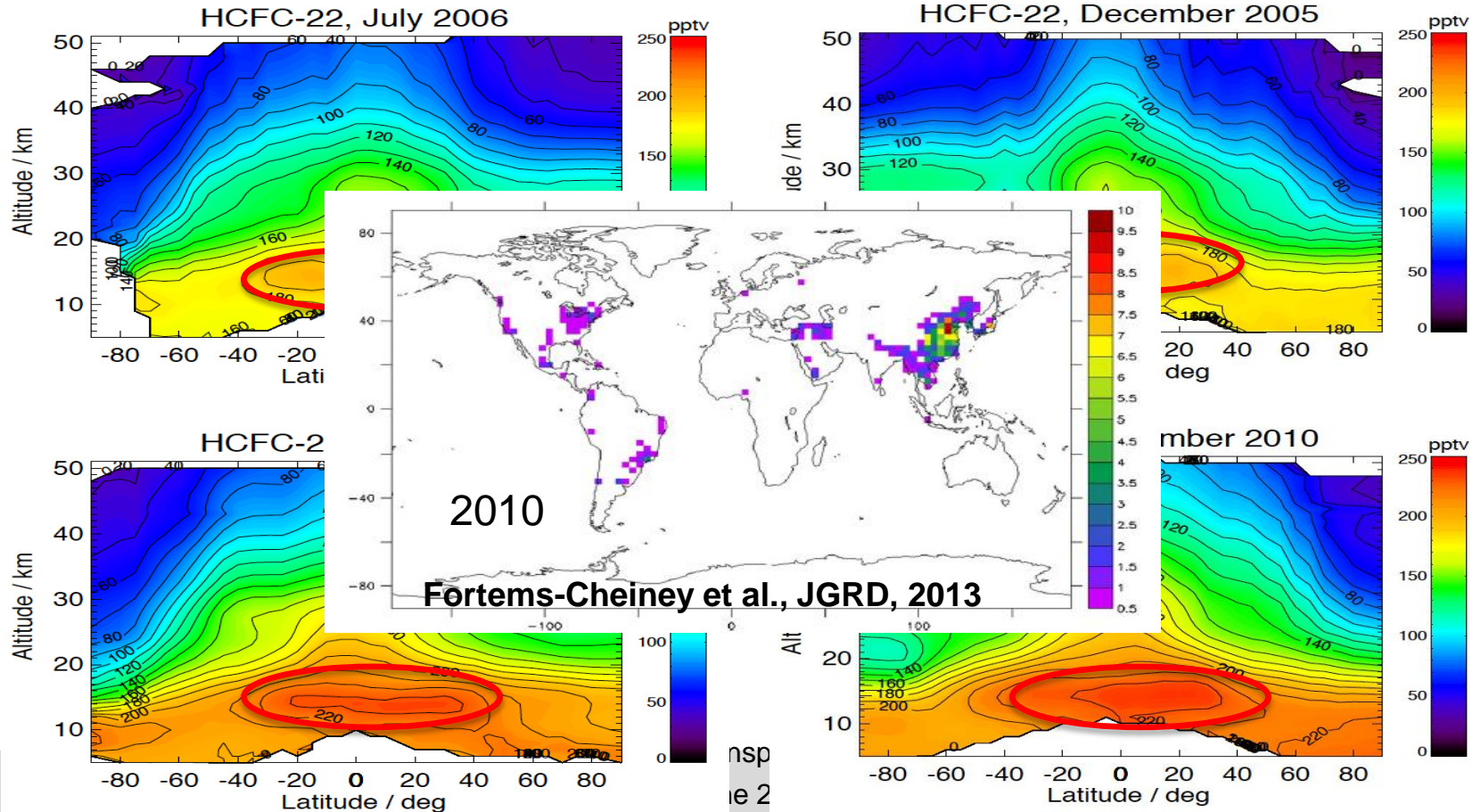


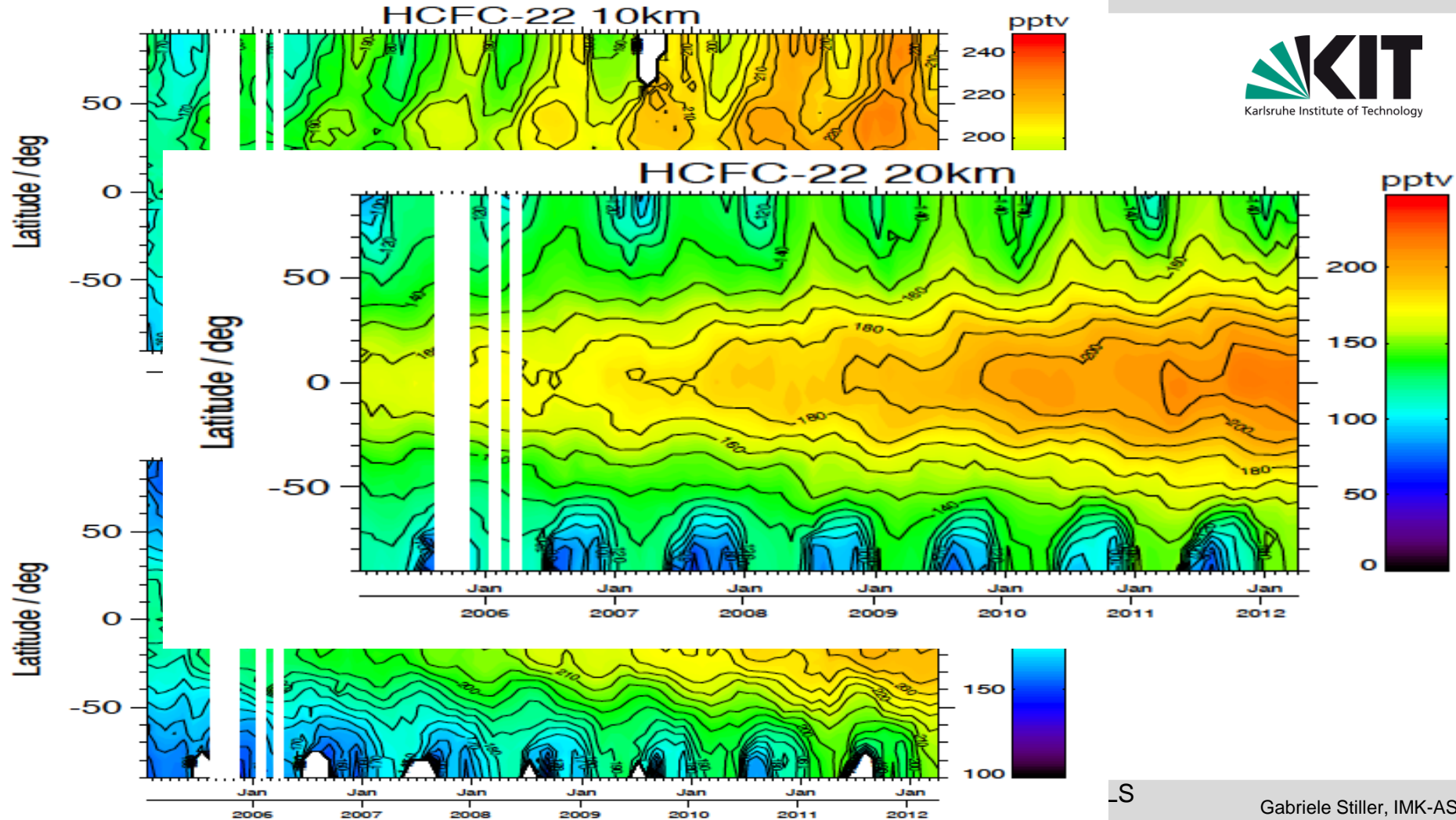
adjustable 0 - 500 km

Is the Asian monsoon system a relevant pathway for direct trop-strat transport?



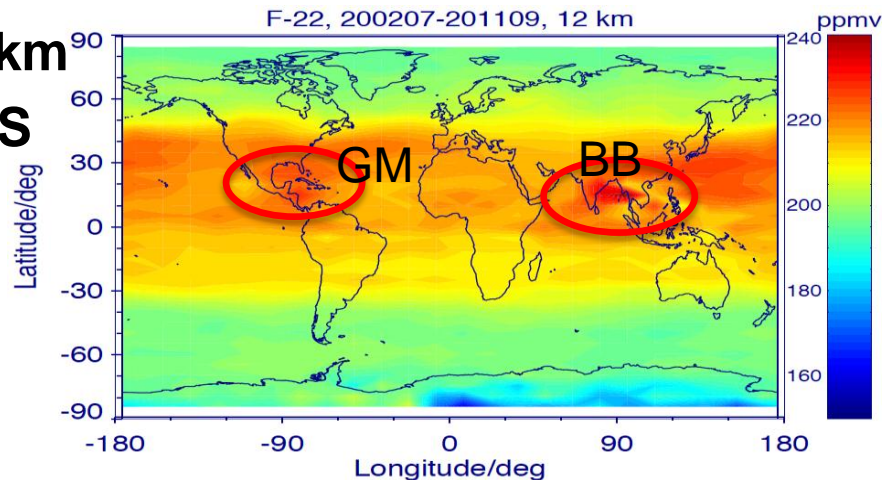
Transport of Asian emissions by the Asian summer monsoon



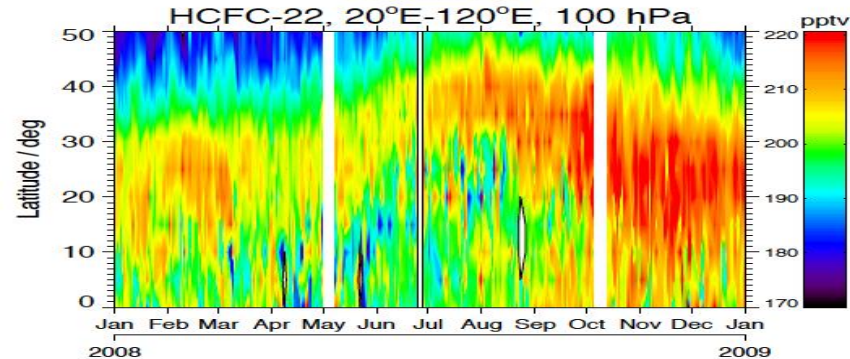
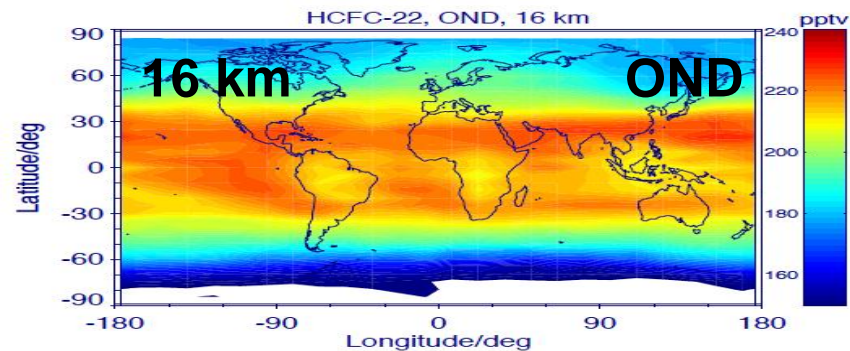
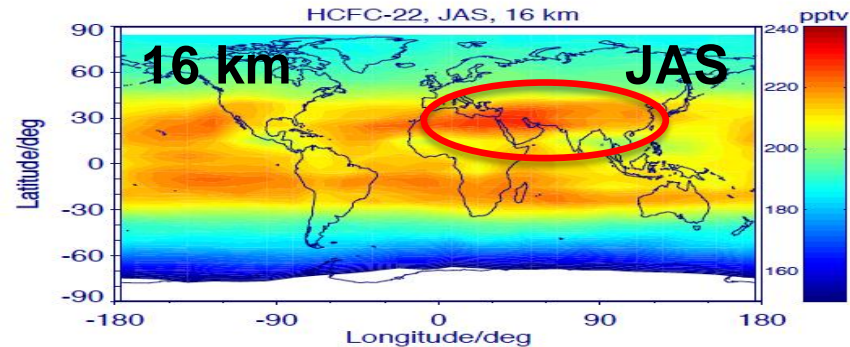
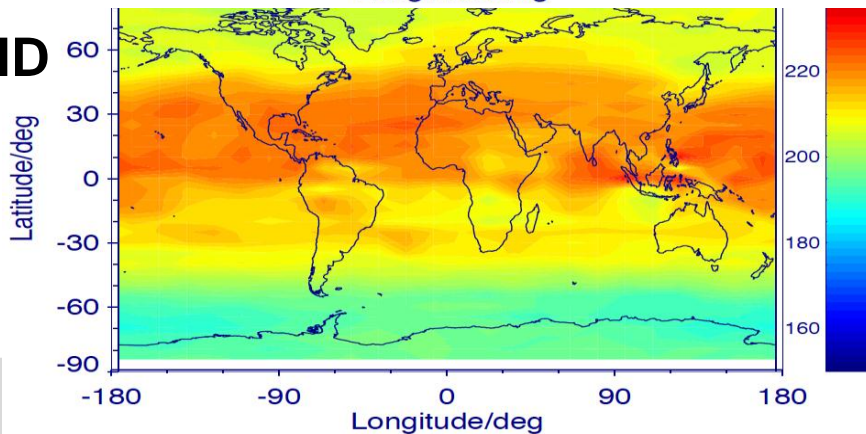


Global maps: 12 and 16 km, JAS and OND

12 km
JAS

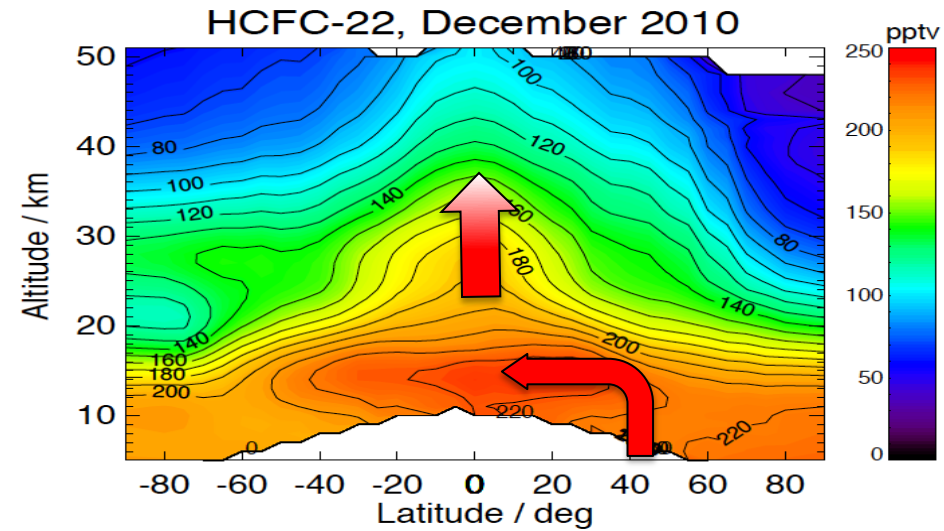
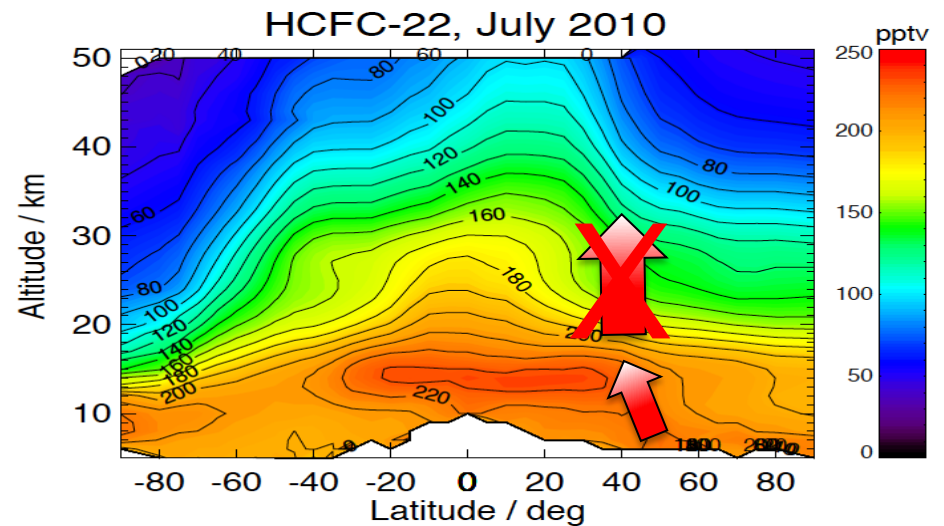


OND

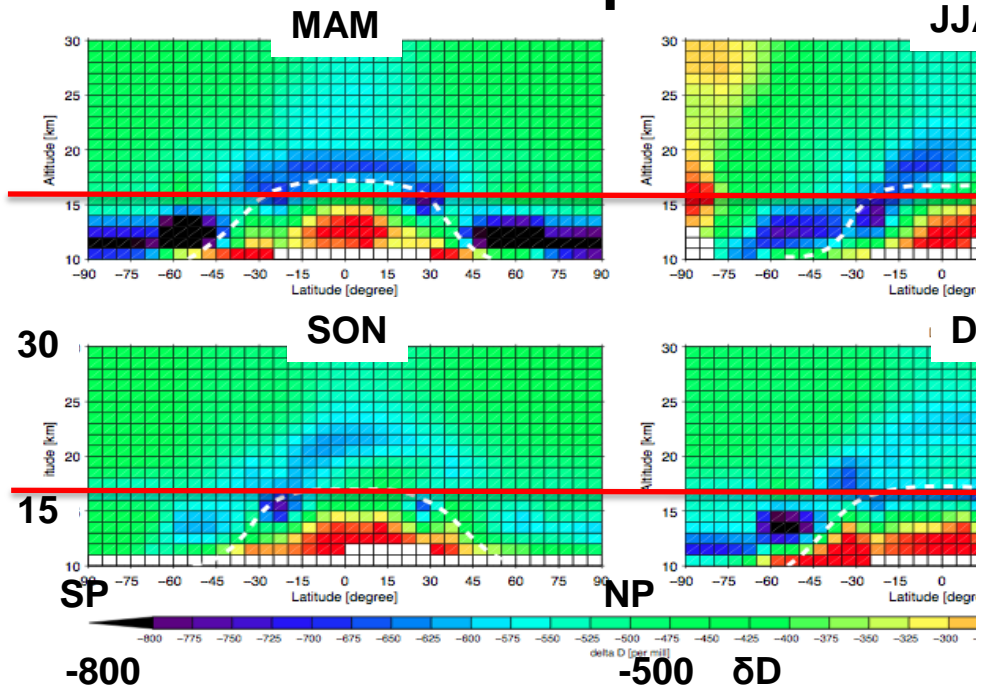


Transport of Asian emissions by the Asian summer monsoon

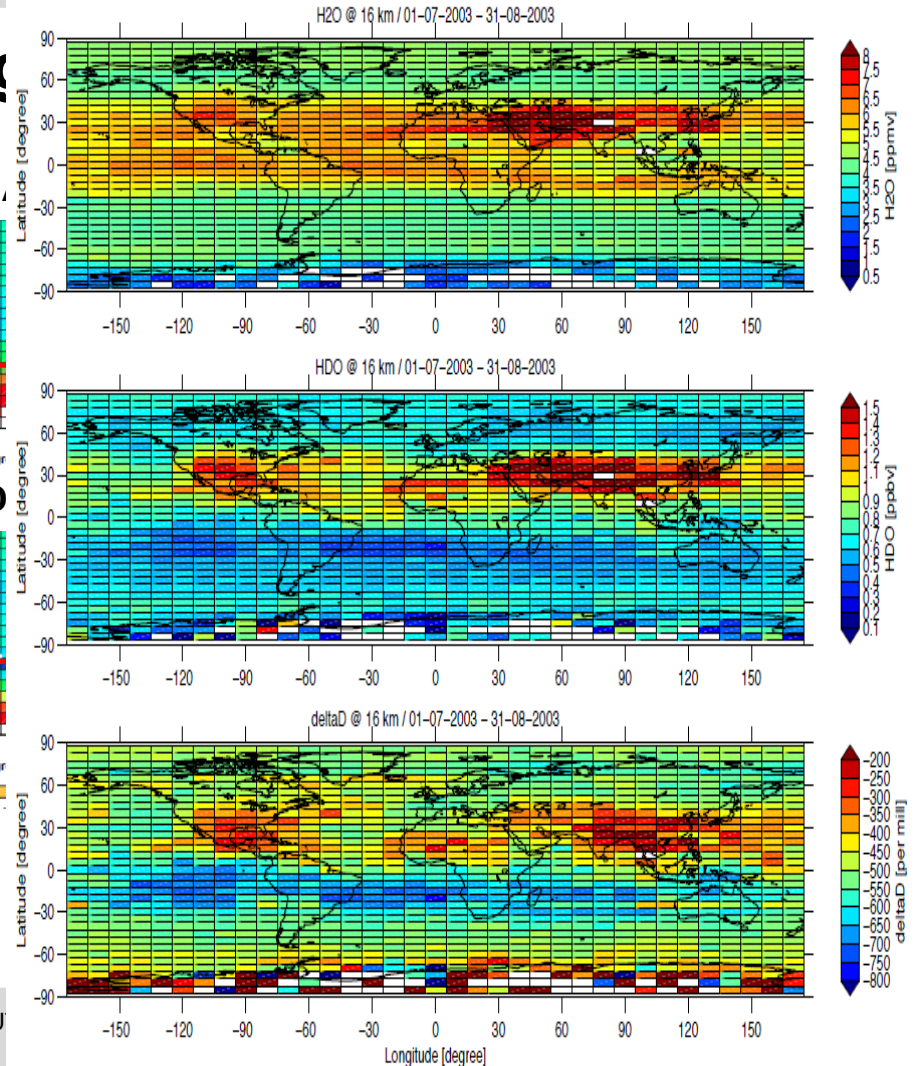
- HCFC-22 abundances of > 240 pptv are higher than observed at any ground-based station
- East-Asian emissions are uplifted in the monsoon convective systems and trapped within the AMA during summer
- They are distributed **below the tropopause** over the entire tropics
- They are uplifted in the tropical pipe to the stratosphere
- No significant direct TSE from the AMA in the subtropics



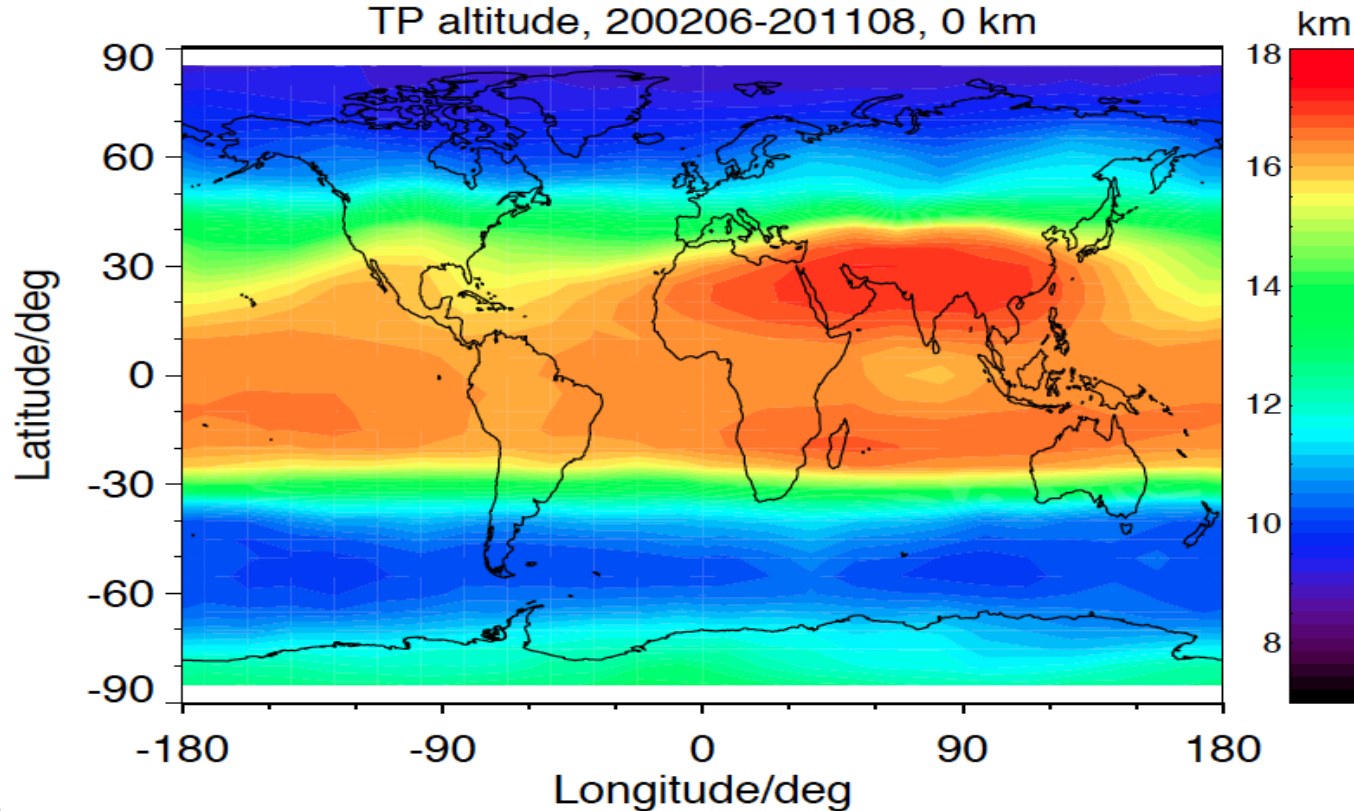
Transport of water vapor through the troposphere: monsoon impact



Lossow et al, in preparation

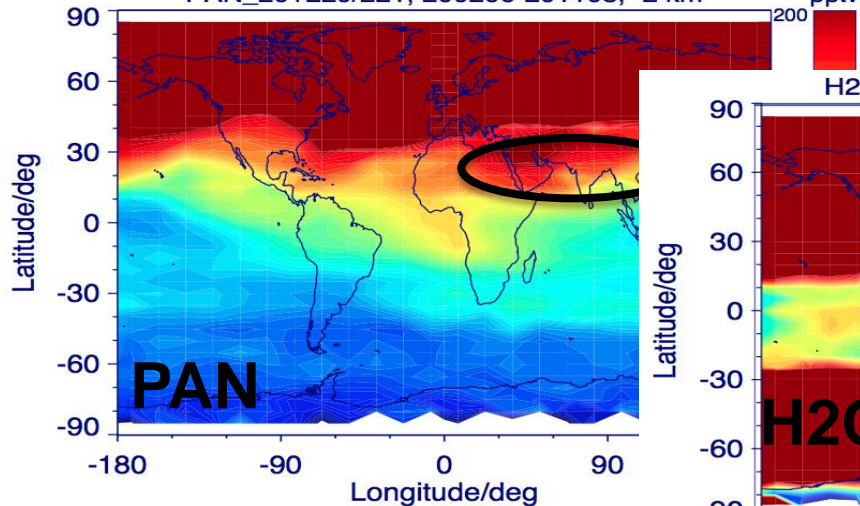


Tropopause height for JJA (from MIPAS temperature profiles)

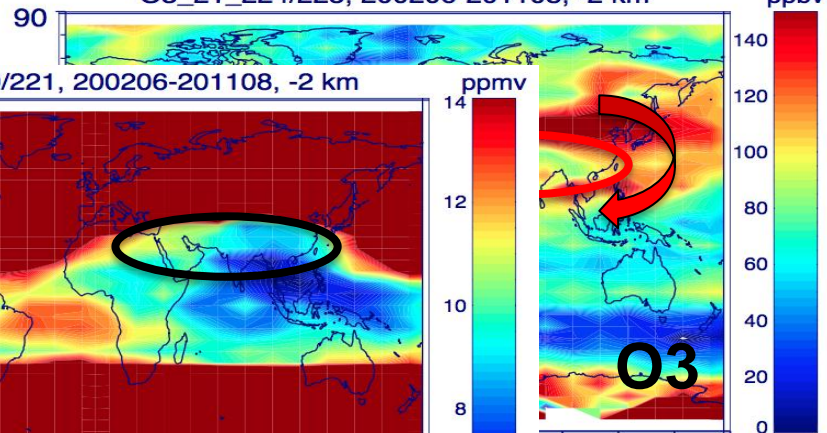


2 km below the tropopause

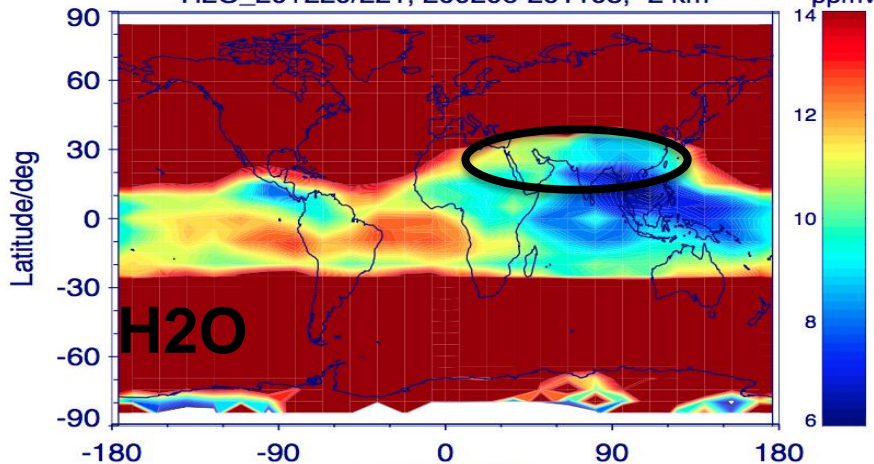
PAN_20+220/221, 200206-201108, -2 km



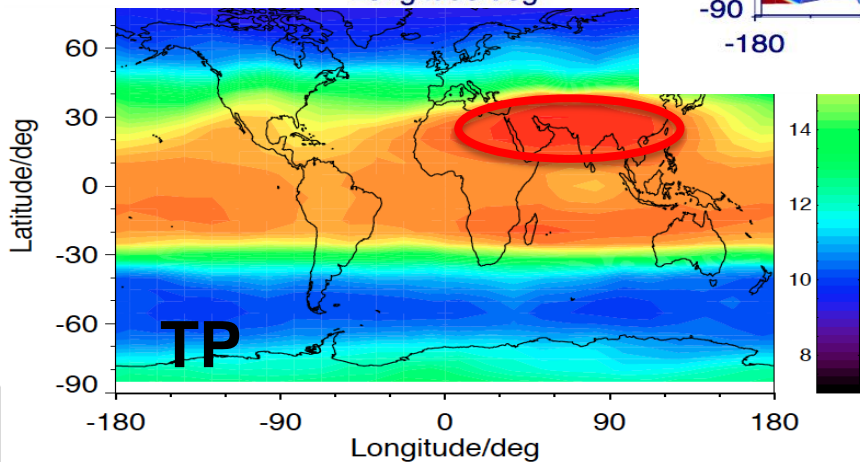
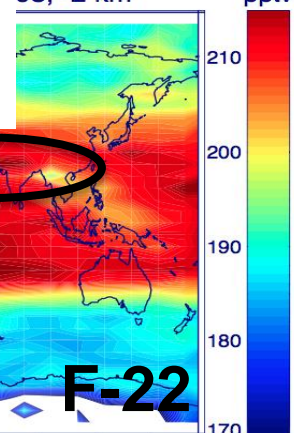
O3_21_224/225, 200206-201108, -2 km



H2O_20+220/221, 200206-201108, -2 km

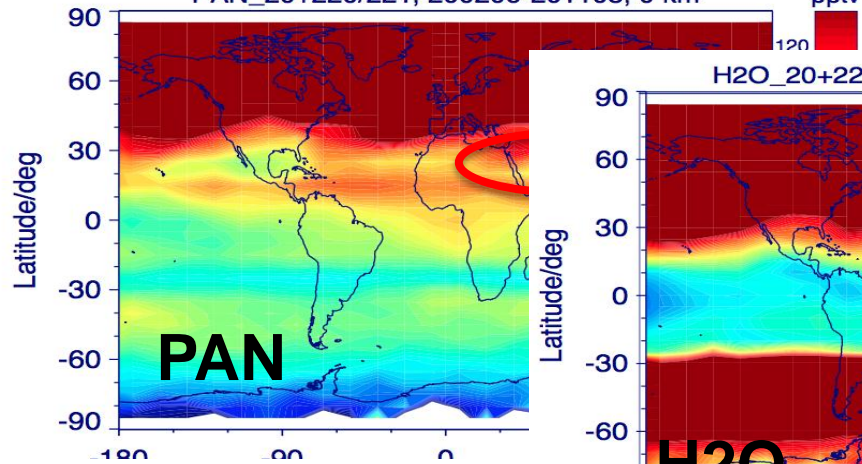


08, -2 km

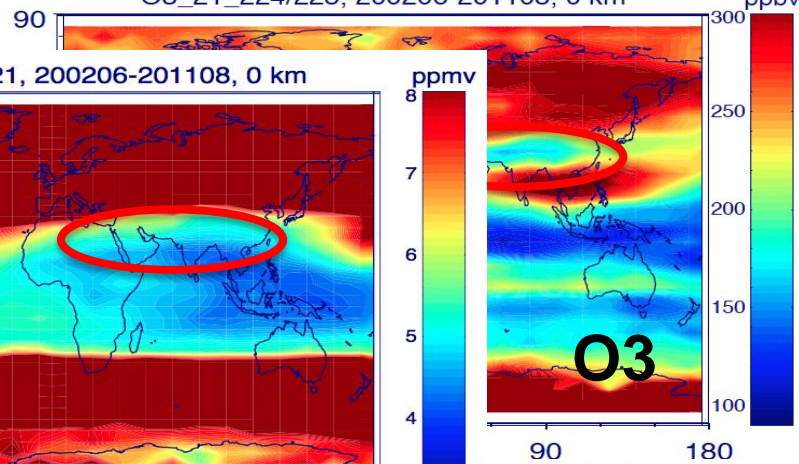


At the tropopause

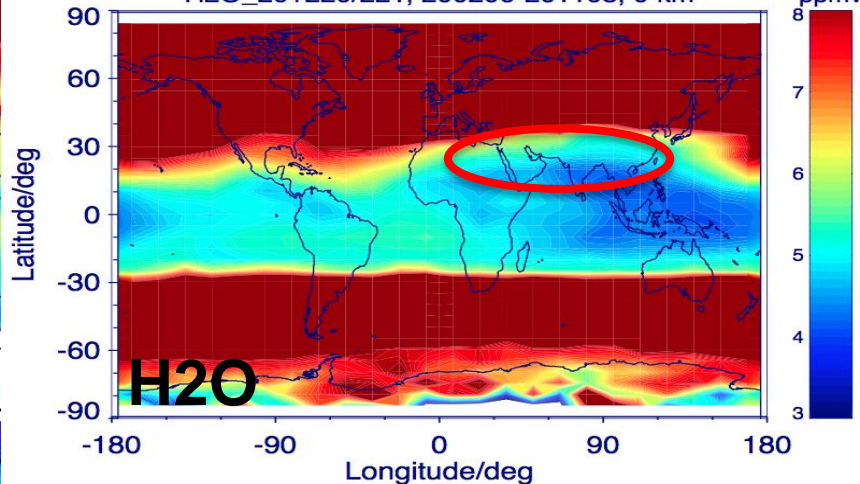
PAN_20+220/221, 200206-201108, 0 km



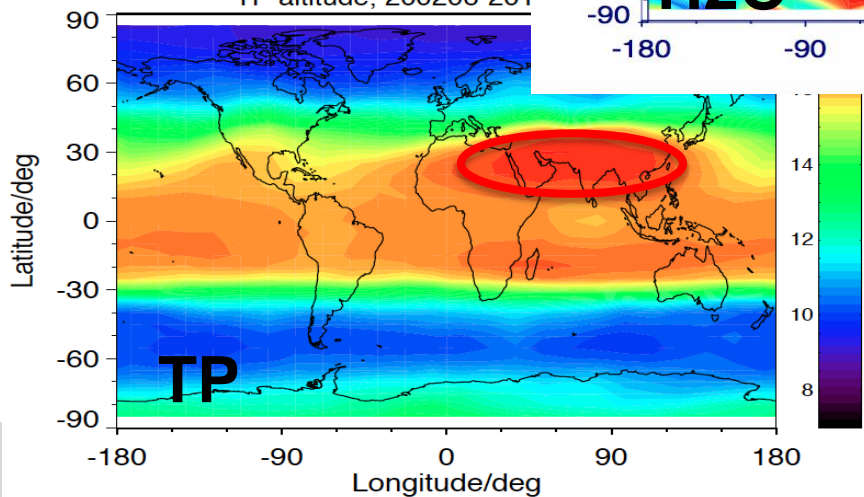
O3_21_224/225, 200206-201108, 0 km



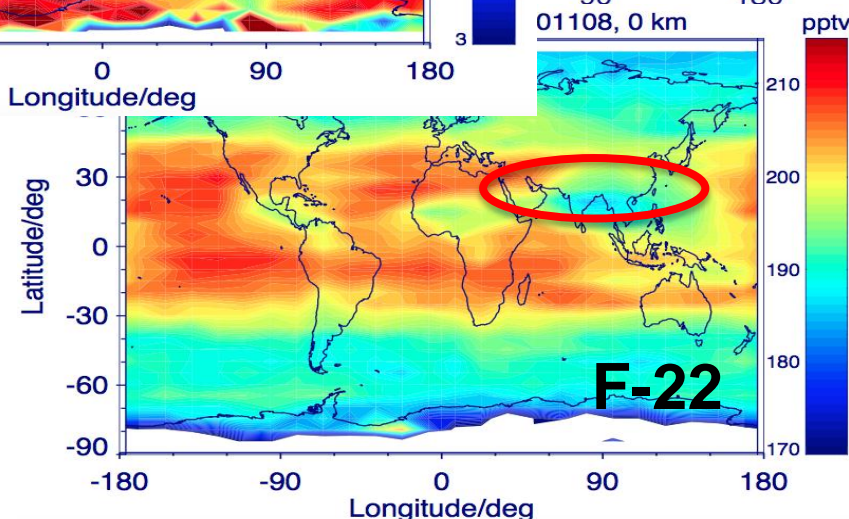
H2O_20+220/221, 200206-201108, 0 km



TP altitude, 200206-201

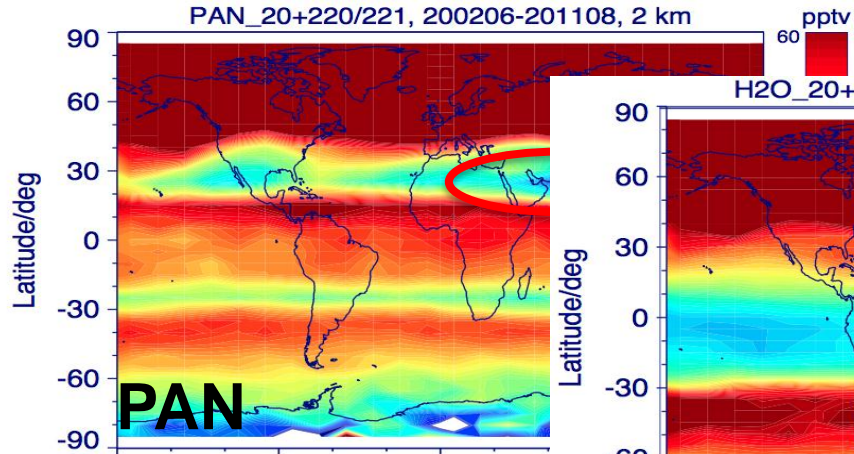


Longitude/deg

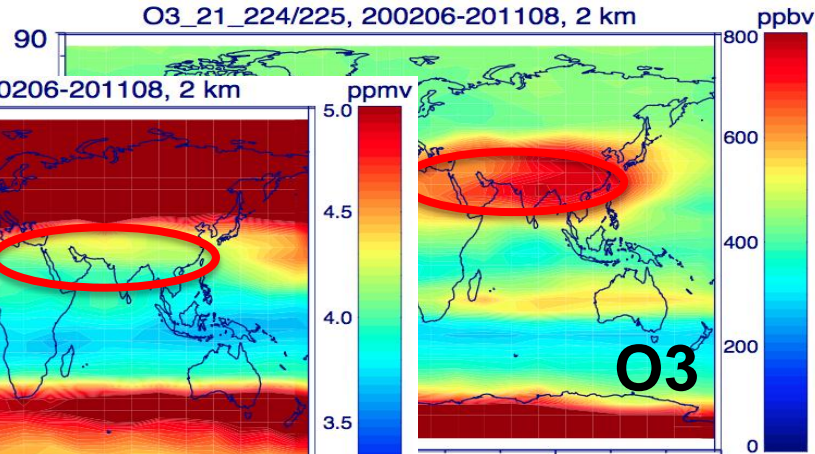


2 km above the tropopause

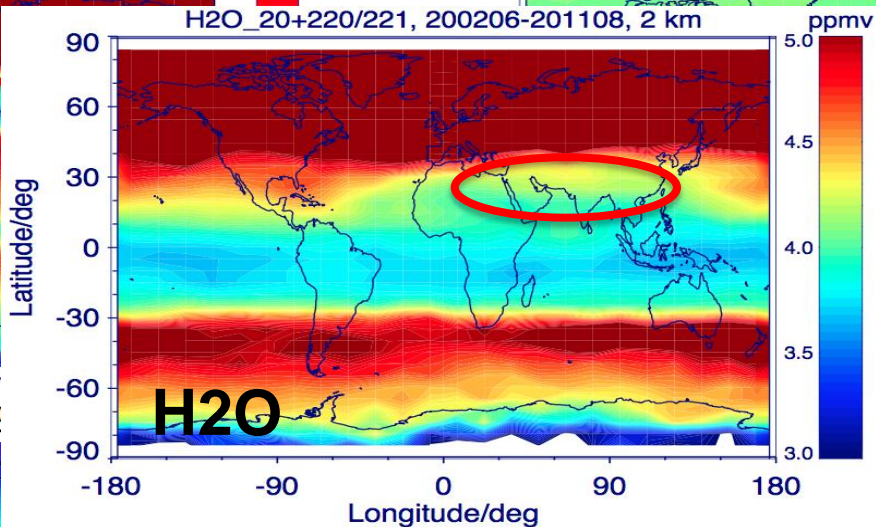
PAN_20+220/221, 200206-201108, 2 km



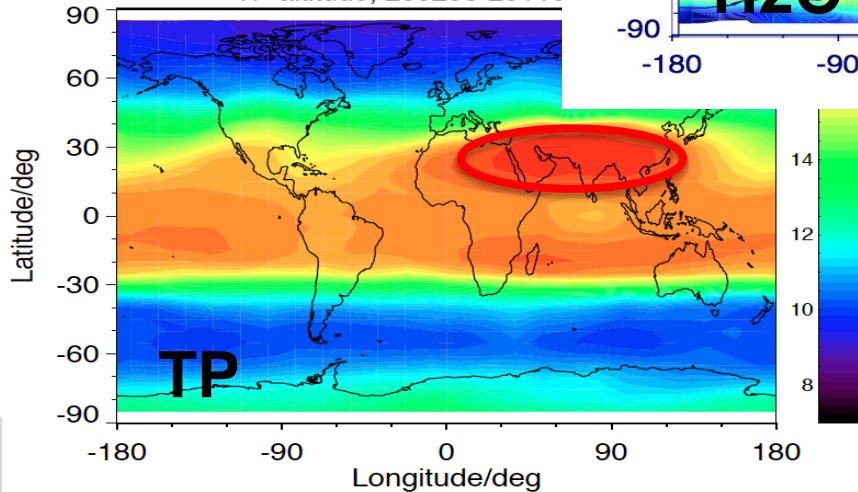
O3_21_224/225, 200206-201108, 2 km



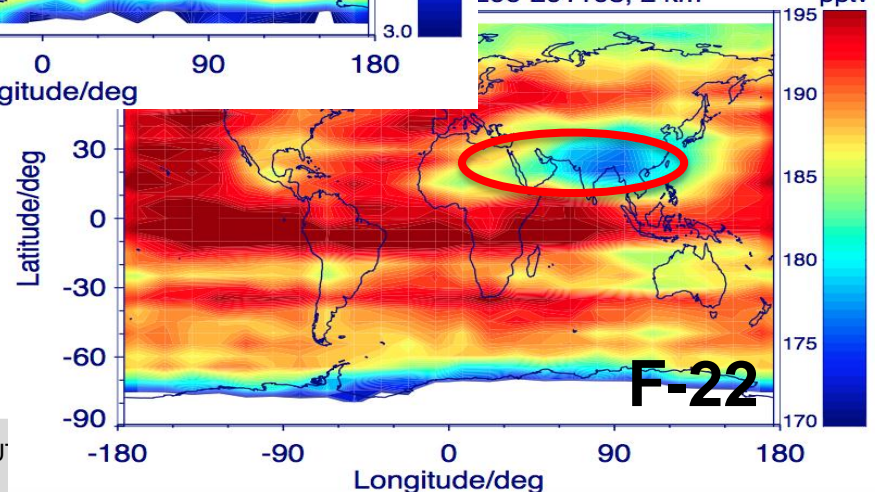
H2O_20+220/221, 200206-201108, 2 km



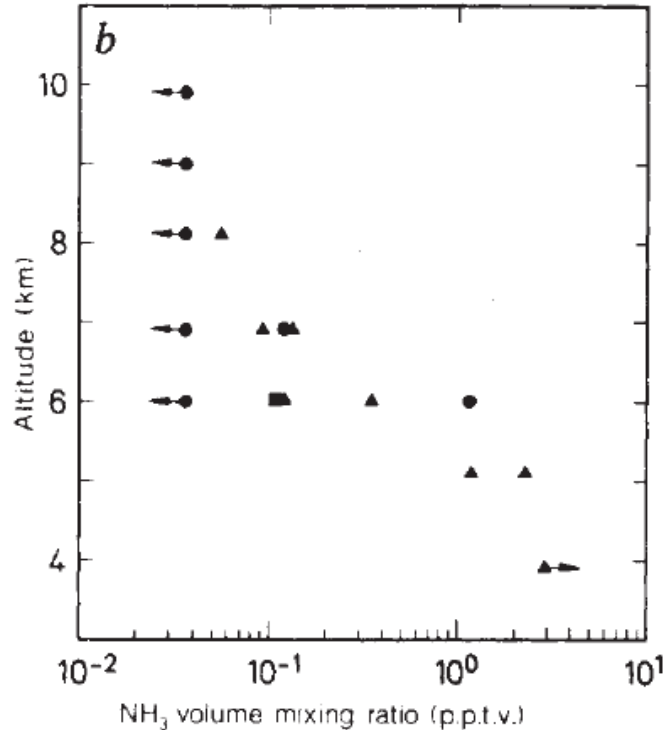
TP altitude, 200206-201108



206-201108, 2 km



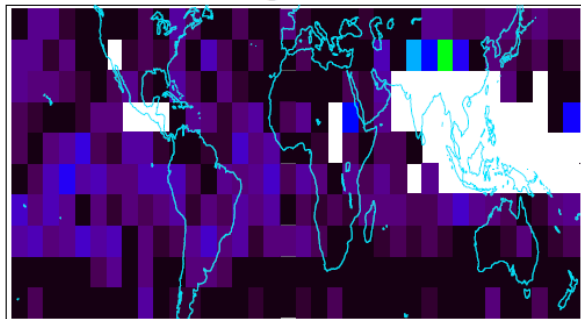
Ammonia (NH_3) in the UTLs?



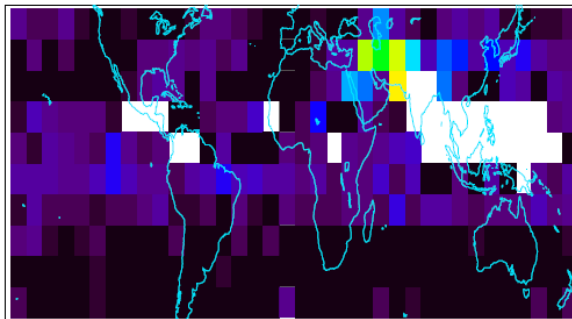
- Only measurements of ammonia in the UT by Ziereis and Arnold, Nature, 1986
- NH_3 volume mixing ratios clearly below 0.1 ppt in the UT

NH₃ distribution 50N/S: Jun/Jul/Aug 2005-2011

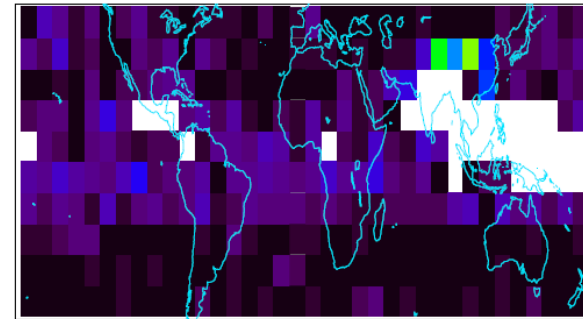
Jun/Jul/Aug 2007 15.0 km



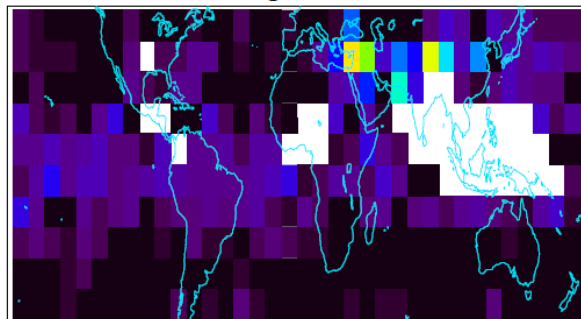
Jun/Jul/Aug 2008 15.0 km



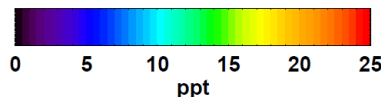
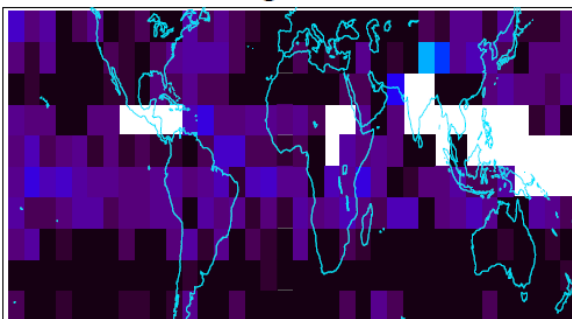
Jun/Jul/Aug 2009 15.0 km



Jun/Jul/Aug 2010 15.0 km

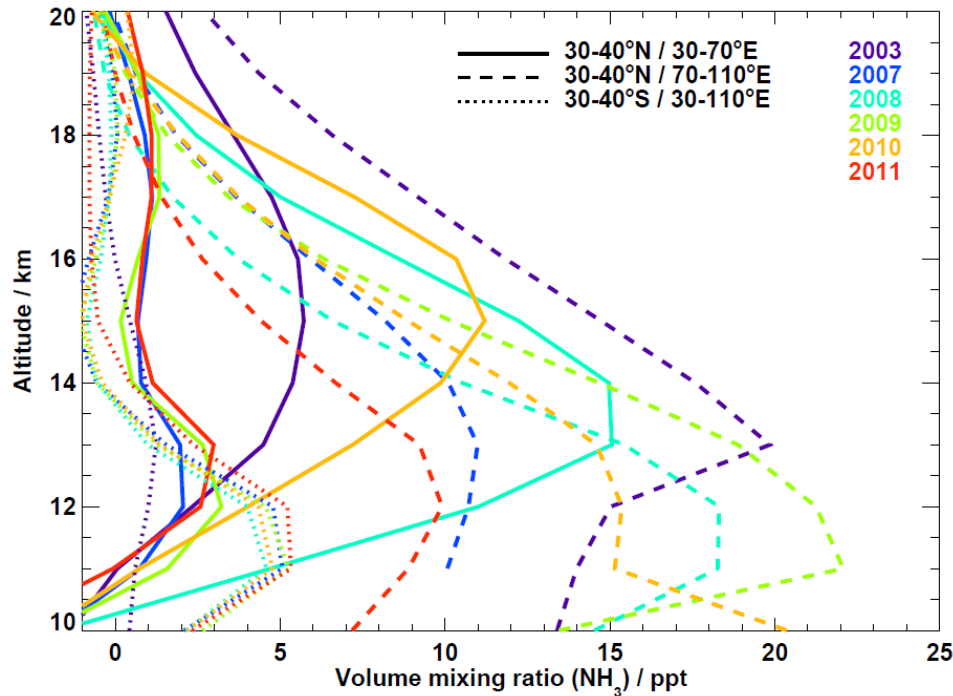


Jun/Jul/Aug 2011 15.0 km



- Total retrieval error
~ 5 ppt
- Random noise error
~1 ppt

NH₃ profile comparisons



- NH₃ maximum larger and at lower altitudes in the eastern part of the monsoon area - - - -
- Much more variable and peaking at higher altitudes in the western part -----
- Maximum in southern hemisphere indicates detection limit of ~5 ppt

NH₃ and the ATAL?

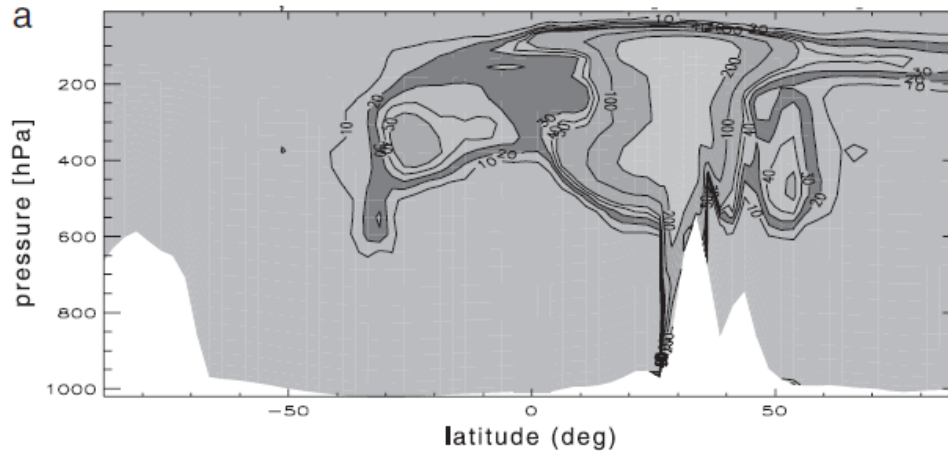
Metzger et al., 2002

Gas/aerosol partitioning

2. Global modeling results

[24] Figures 4a and 4b illustrate that ammonium nitrate aerosol may be formed and transported to the free troposphere, even in summer. Remarkably, concentrations can reach up to 200 pptv at 200 hPa over the convective region south of the Himalaya (Figure 4a, top). The only possibility to form nitrate in our model is through aerosol neutralization by ammonia. In Europe during summer there is not

Aerosol Nitrate, August 1997 (monthly mean) [lon = 80°E]



Europe, which actually originates from southern Asia. The explanation is that in northern India ammonia is emitted at relatively high altitudes. During the Indian summer strong monsoon convection is associated with upper tropospheric dry easterlies, while substantial amounts of ammonia are convectively transported above the boundary layer. The convective clouds through which the ammonia is transported are hardly acidic so that the ammonia is only partly dissolved and removed by precipitation. Once transported to the upper troposphere surplus ammonia neutralizes nitric acid that is present in higher amounts than sulfuric acid. As a result, an extended plume of ammonium nitrate is predicted to exist in the upper troposphere. Hence surface emitted ammonia from northern India forms ammonium nitrate that can be efficiently transported in the subtropical jet stream toward Africa and even to southern Europe near the tropopause.

Summary and conclusions



- MIPAS provides many species for the study of the Asian (and American) monsoon UTLS
- HCFC-22 is a good tracer for Indian/Chinese pollution because of its unique sources in East Asia.
- HCFC-22 forms a maximum layer in the tropics below the tropopause, indicating transport from the Asian monsoon UT into the TTL and upward in the tropical pipe.
- Analysis of species at the tropopause: vertical transport through the AMA tropopause rather weak if any; even subsidence above the AMA?
- Strange behaviour of H₂O at the AMA tropopause, more linked to the West Pacific.
- New product: NH₃, enhanced in AMA (for some years); link to ammonium nitrate aerosol and ATAL.

Thank you!