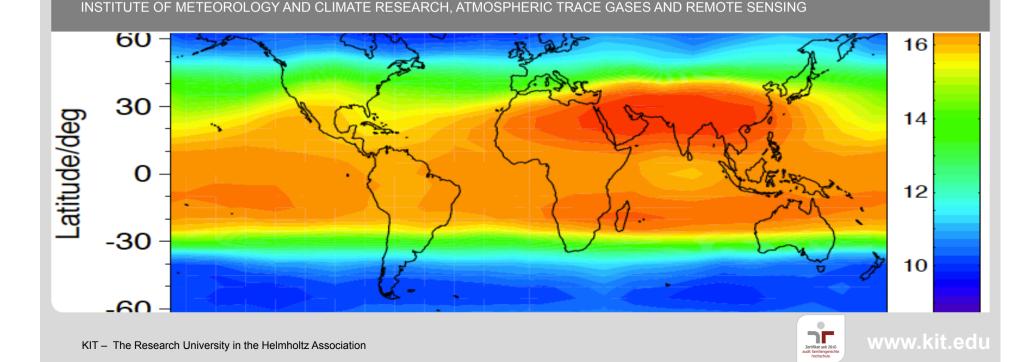


Tropopause-following analysis of water vapour and ΔD for the monsoon systems and the tropics

Gabriele Stiller, Norbert Glatthor, Stefan Lossow, and the MIPAS team at KIT

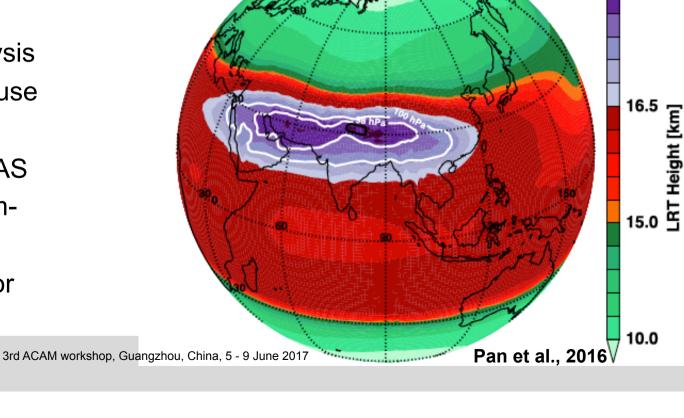


Motivation / open questions

- How is water vapour transported into the stratosphere?
- In particular: What role plays overshooting of convection?
- Are there systematic differences in relevance or frequency for monsoon systems and tropics, respectively?

(a)

- Tropopause altitude varies considerably, on altitude, pressure and potential temperature levels
- Therefore: analysis directly on tropopause surfaces
- Data base: MIPAS observations of temperature, water
 vapour and HDO for
 2002 to 2012



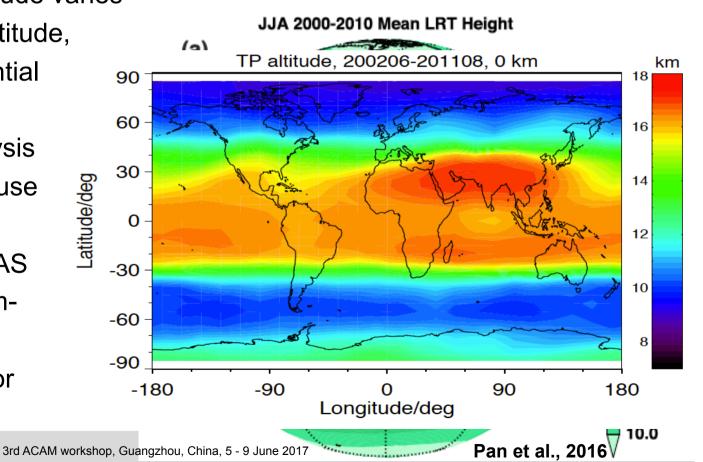
JJA 2000-2010 Mean LRT Height



17.0

Motivation / open questions

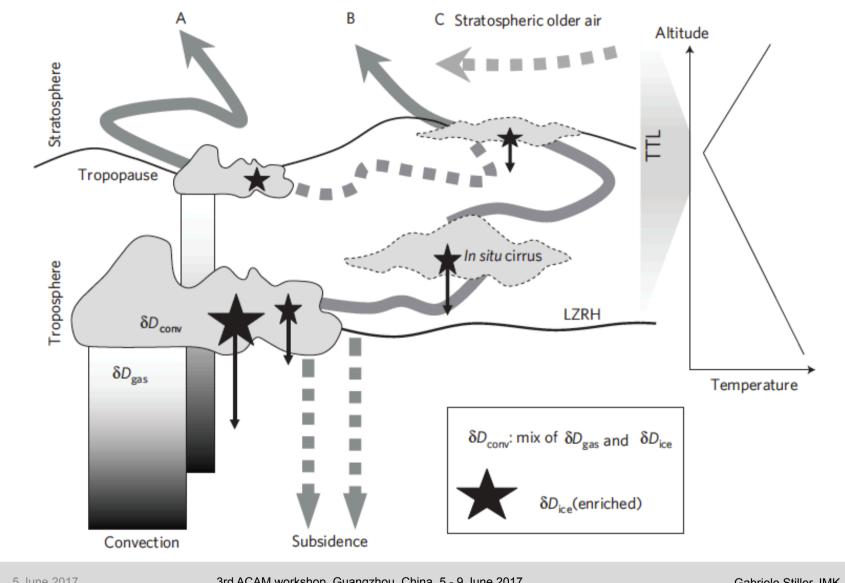
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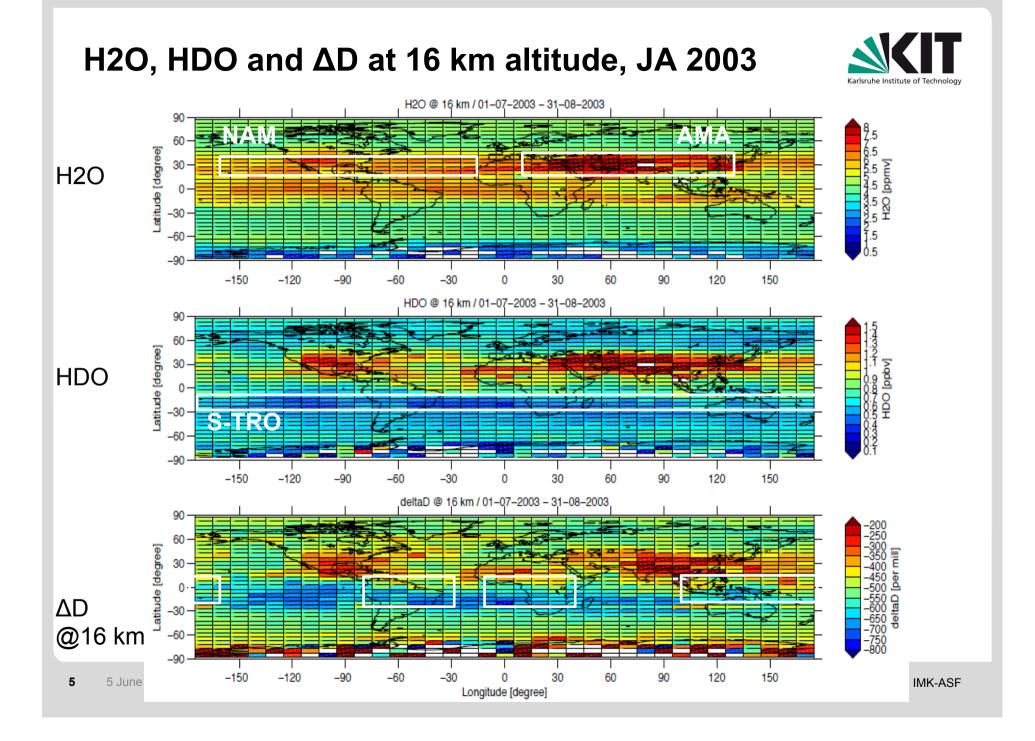


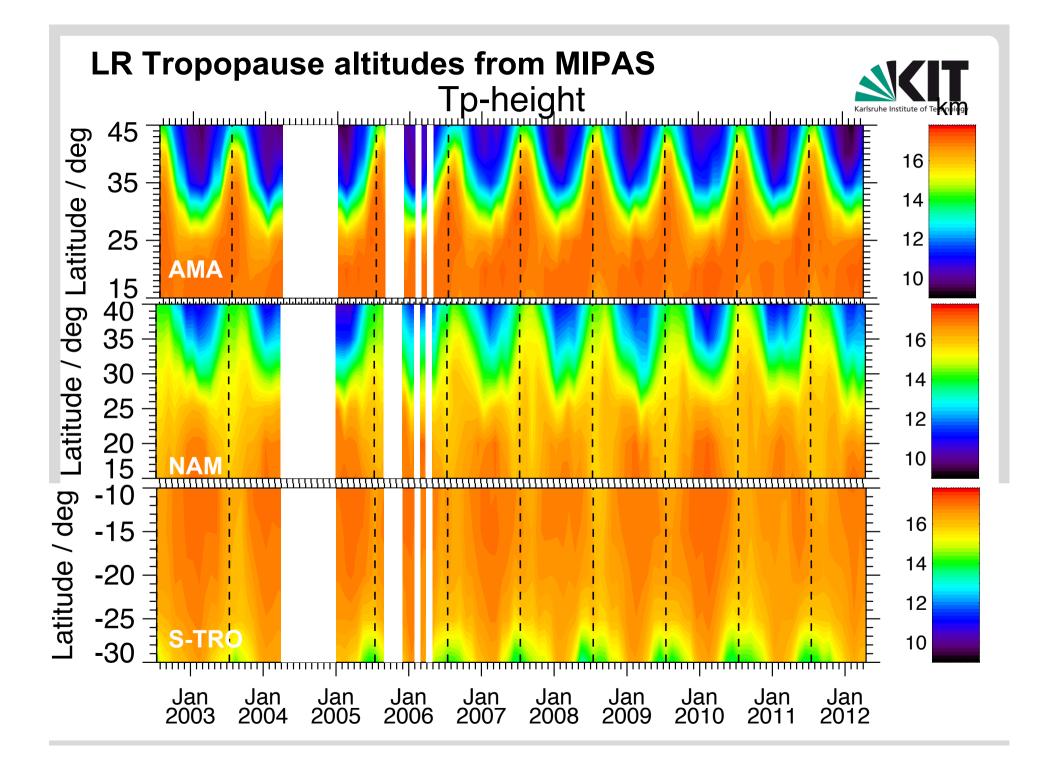


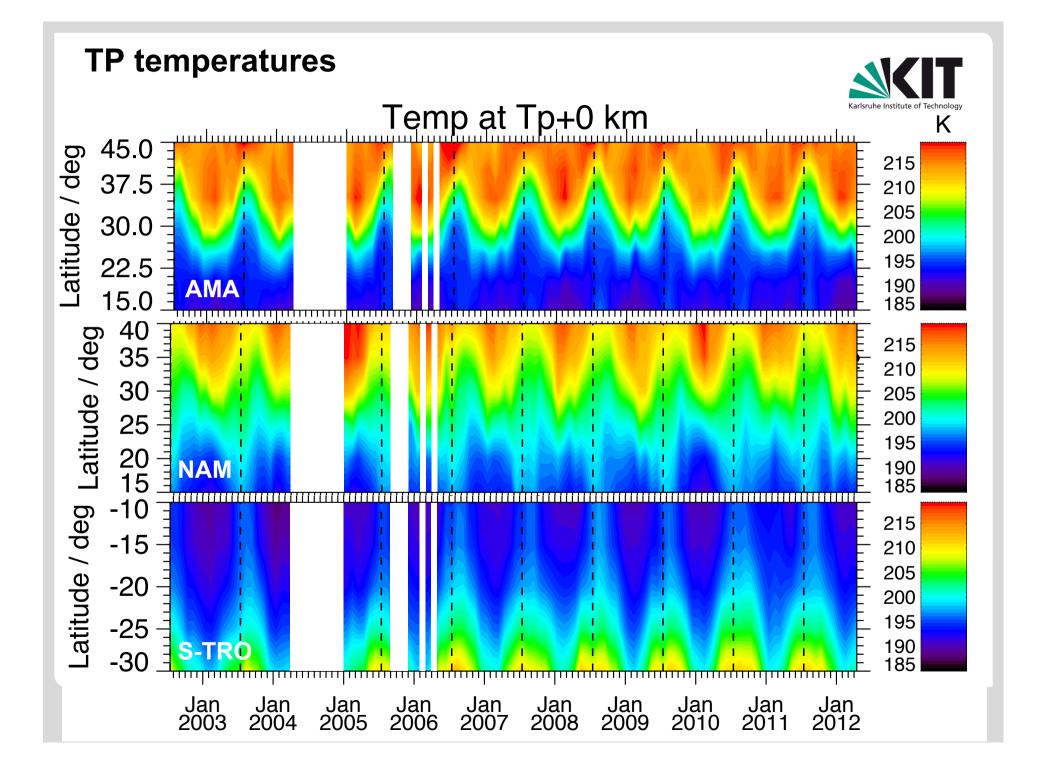
Processes leading to depletion and enrichment of heavy water (ΔD)











General behavior of TP

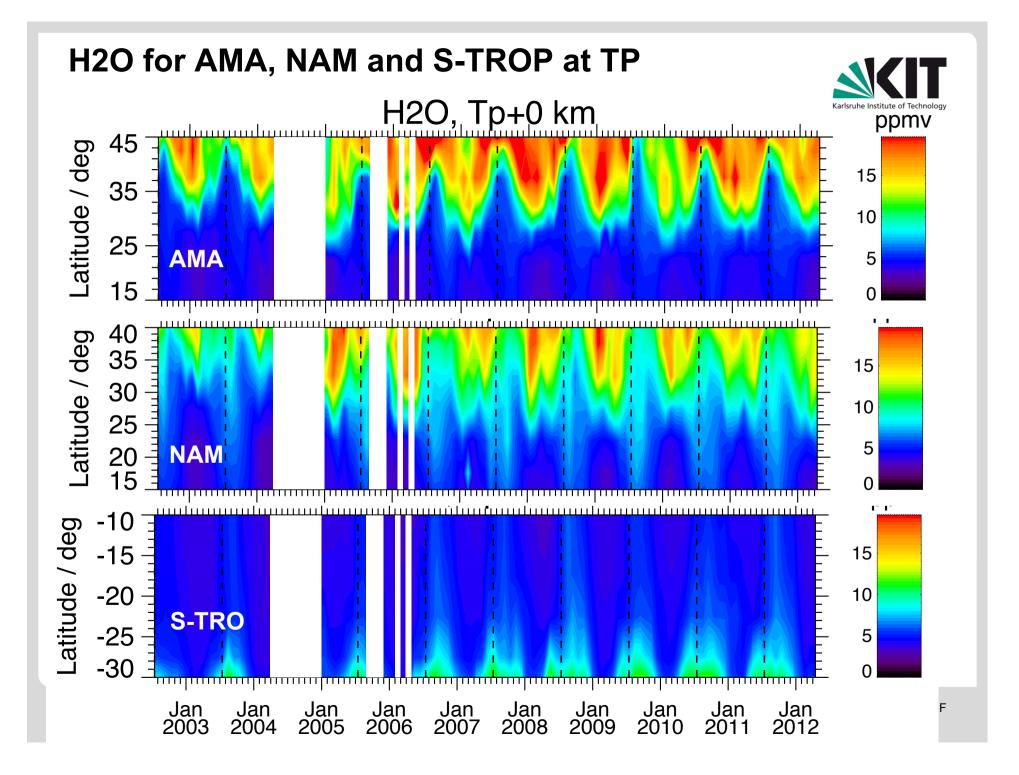


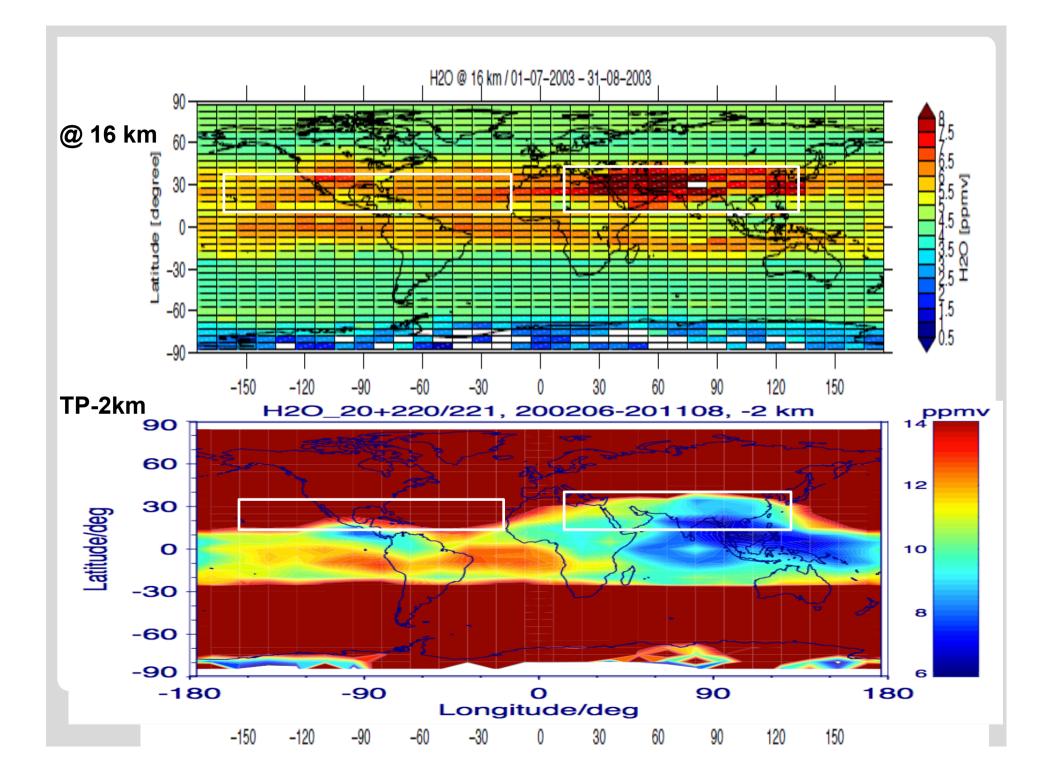
- Altitude:
- S-TRO: TP highest during SH Summer (16-17 km), lowest around July (< 16 km)
- AMA: highest TP on absolute scale, 17-18 km; highest and the most North around 15 July
- NAM: lower than AMA (~15 km), phase and extension as AMA

Temperature:

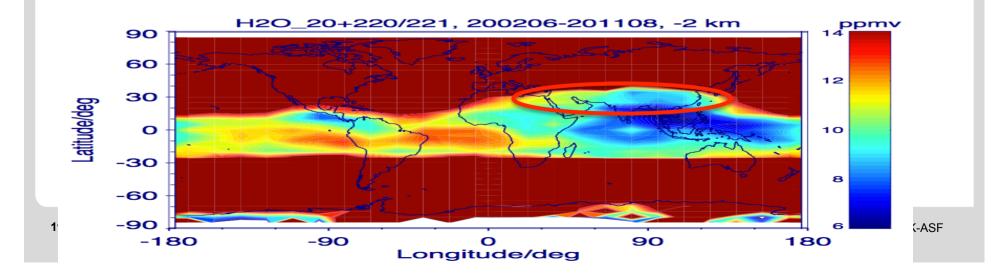
- S-TRO: T lowest during SH summer (<190 K), highest around July (~ 200K); → lowest WV above TP in SH summer; lowest WV absolute
- AMA: coldest and furthest North around July, T ~ 195 K; strong temperature contrast and meridional gradient between monsoon and extra-tropics; → low WV during monsoon season, but higher than tropics
- NAM: less cold than AMA (~200 205 K), less strong meridional gradient; → highest WV of all three regions expected

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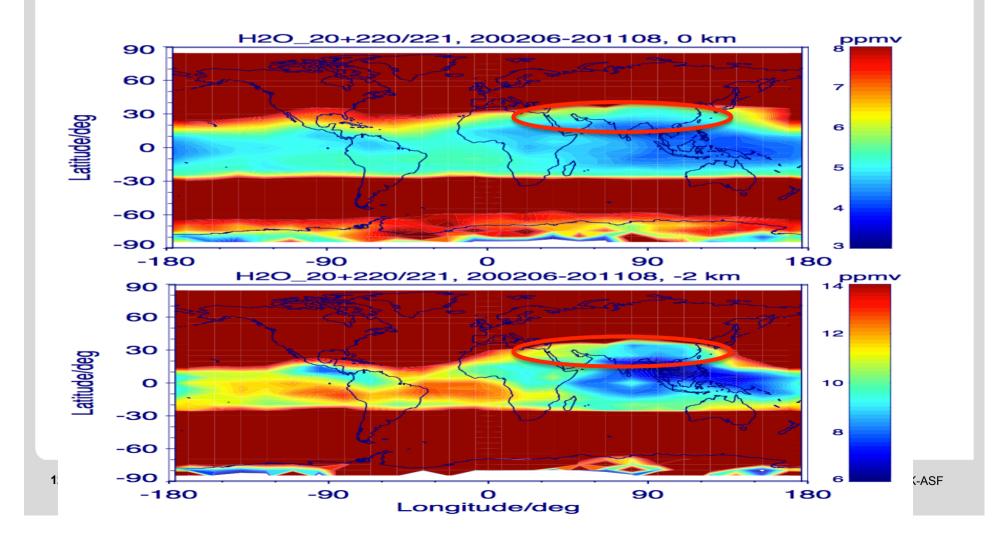


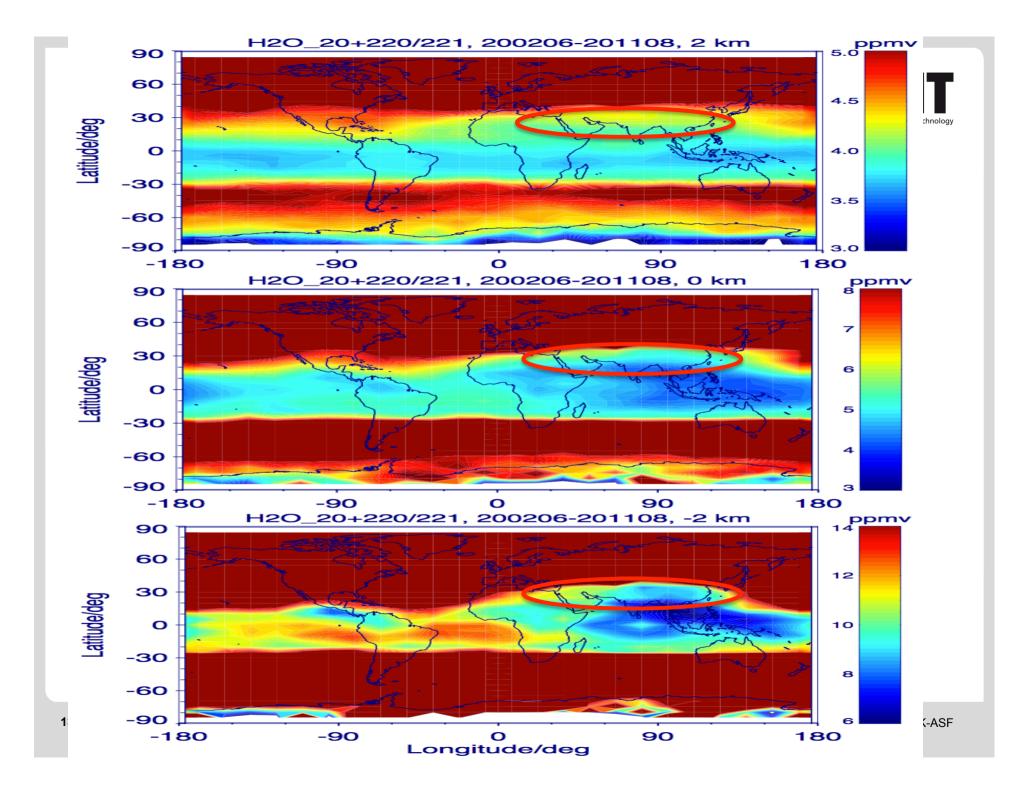


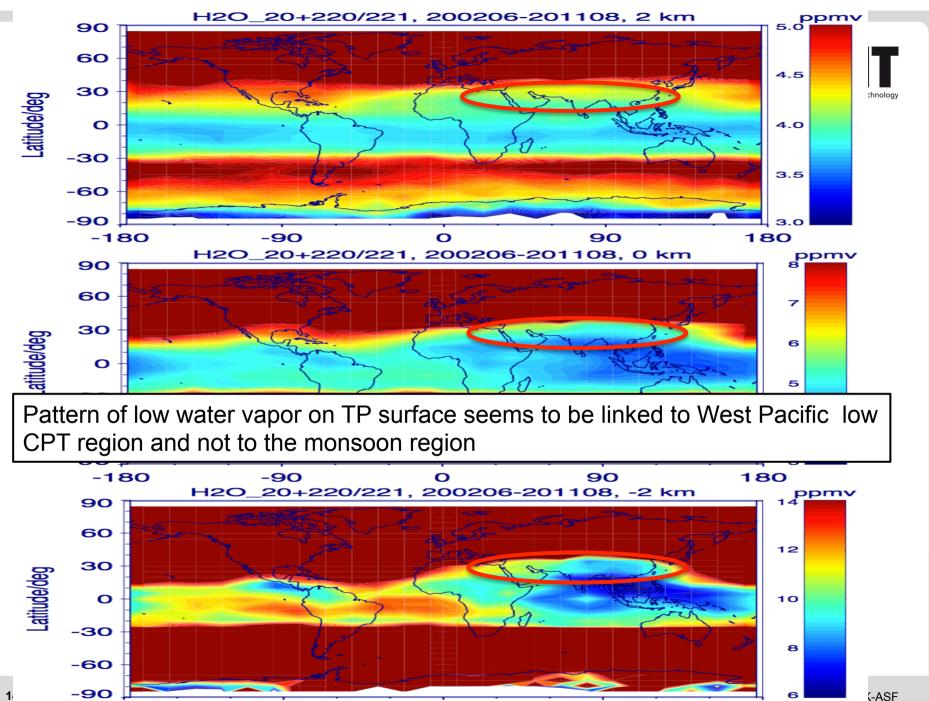












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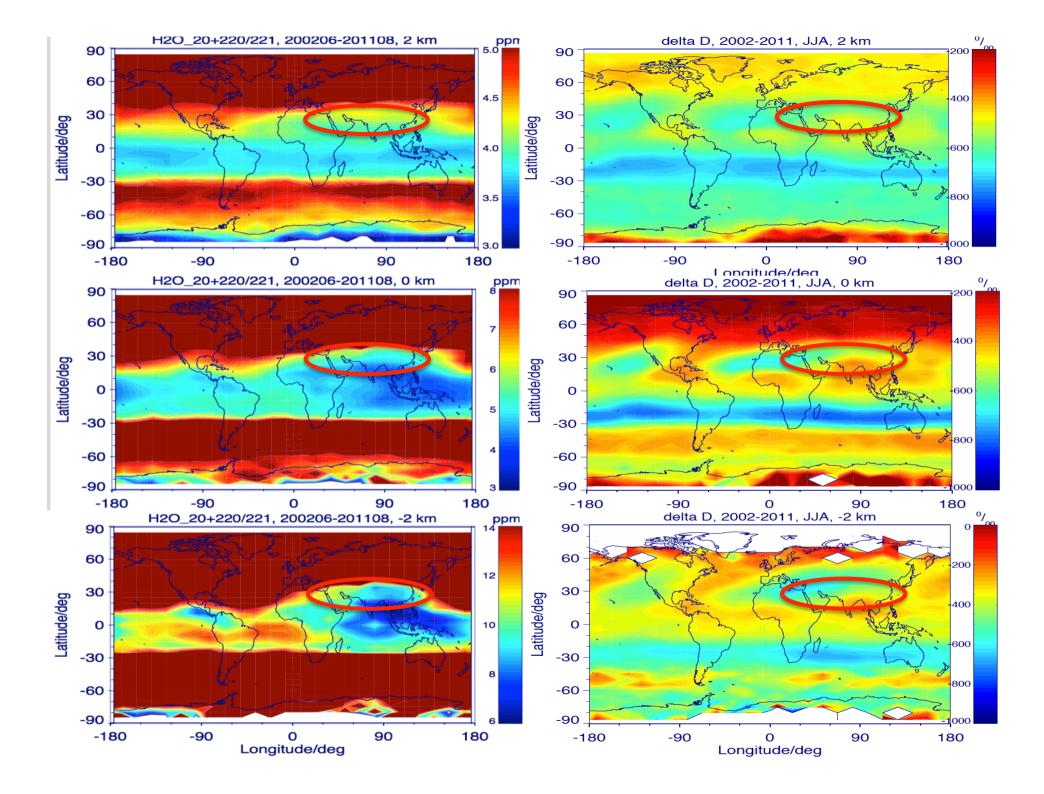
Longitude/deg

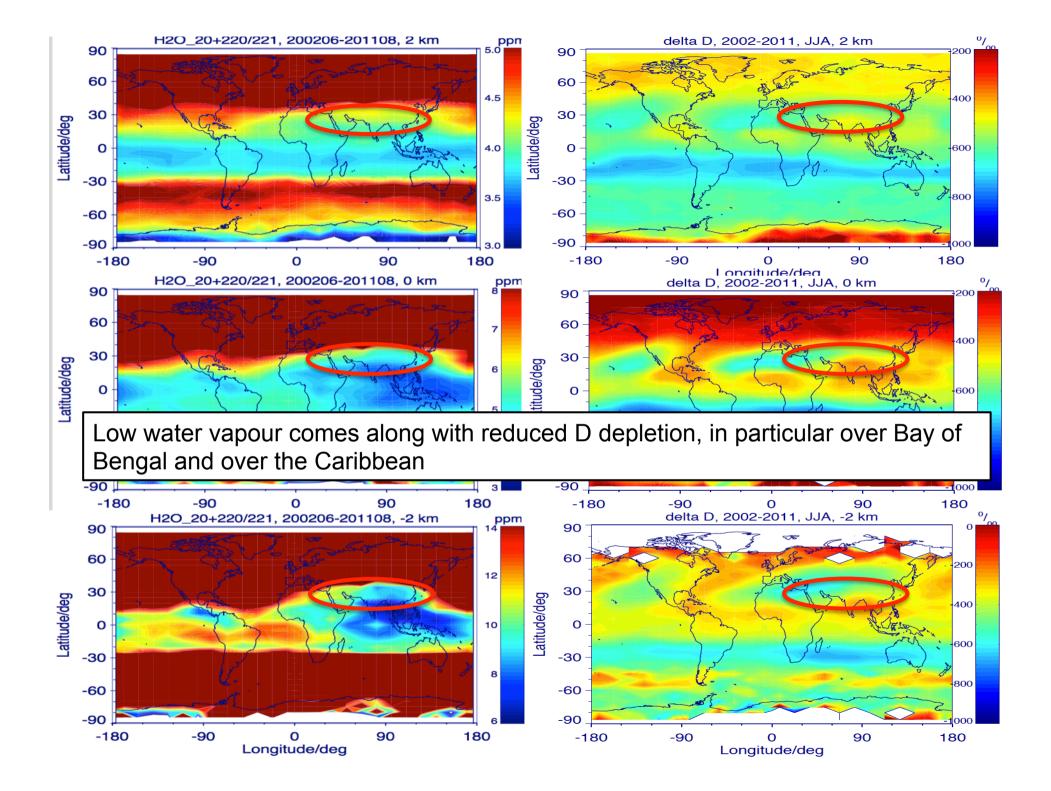
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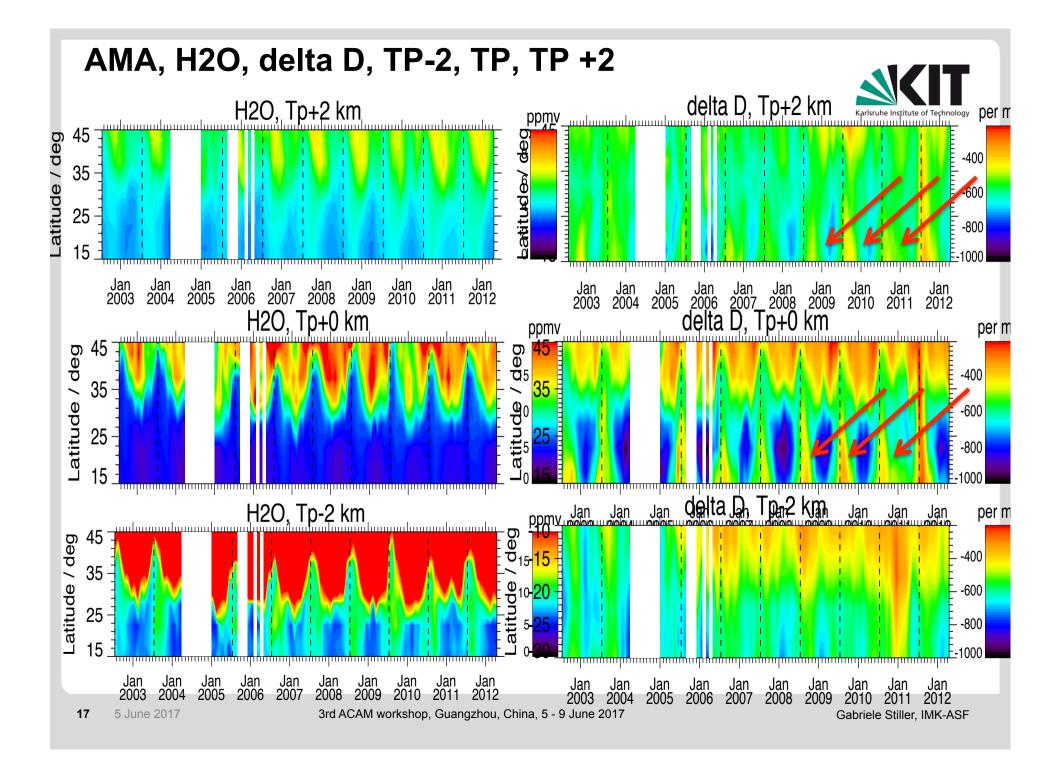
180

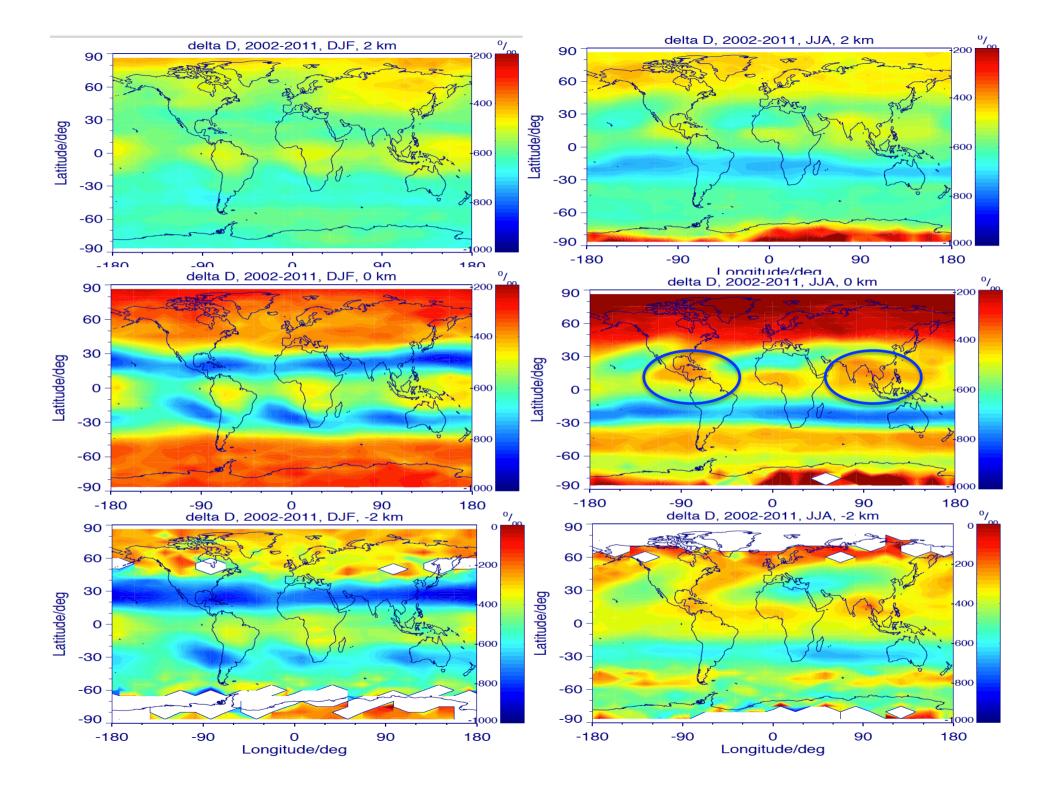
-90

-180





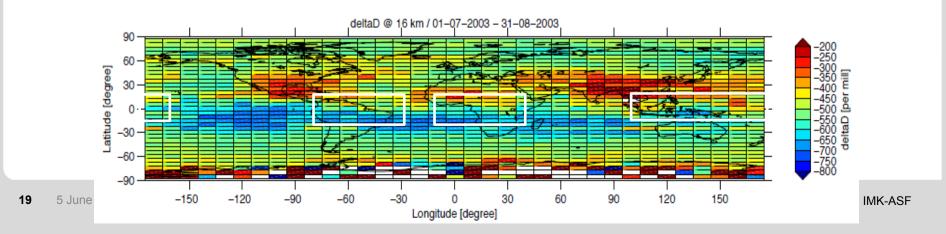


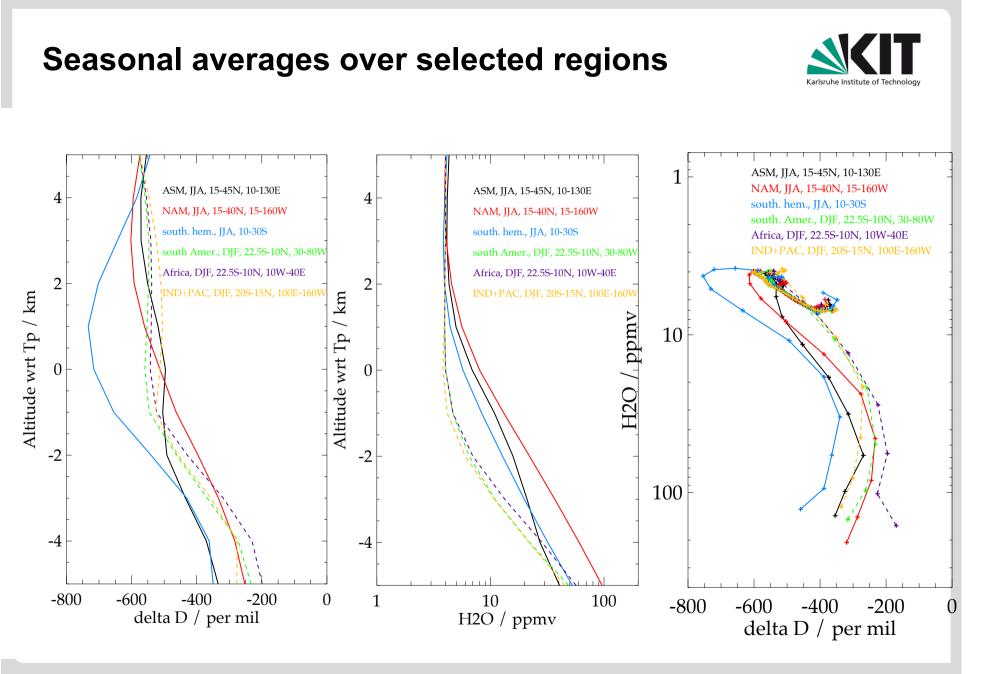


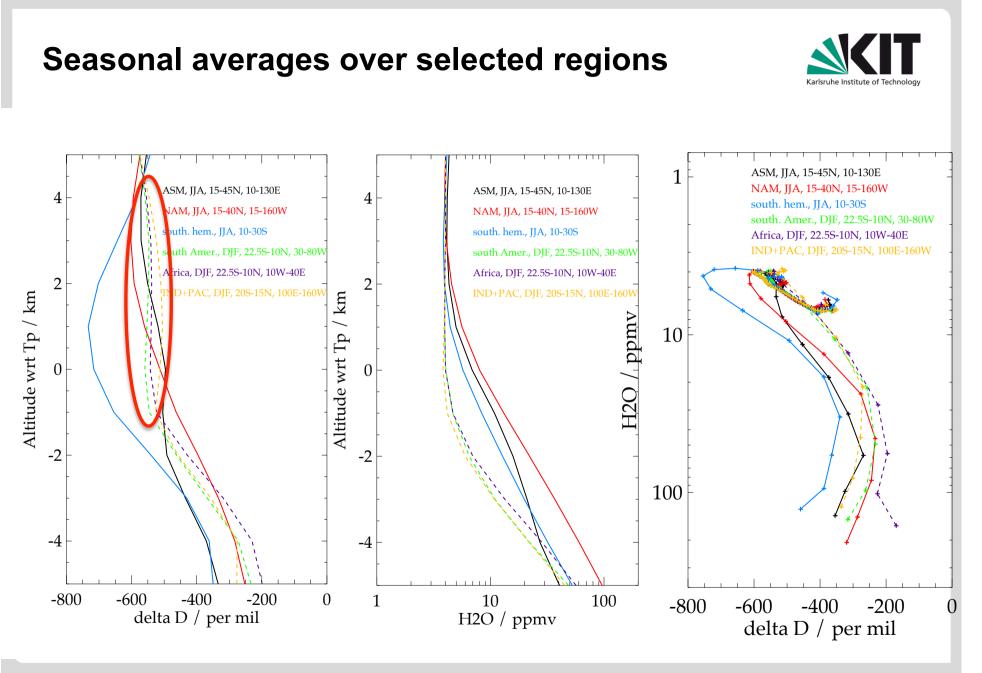
Summary so far

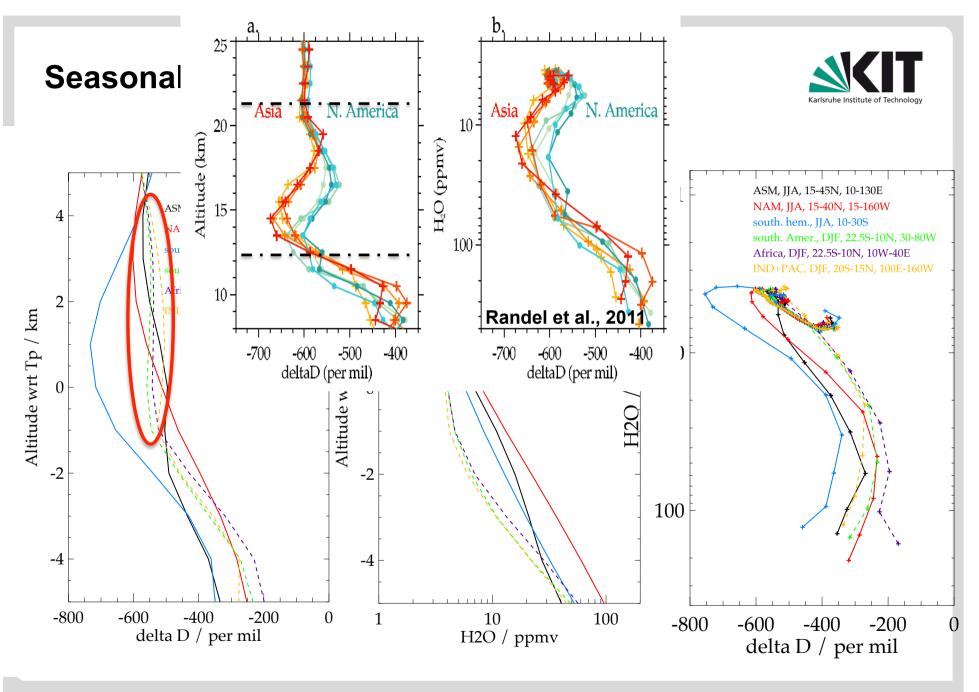


- □ Low water vapor just below and at monsoon tropopause
- Pattern of low water vapor on TP surface seems to be linked to West Pacific low CPT region and not to the monsoon region
- Low water vapour comes along with reduced D depletion, in particular over Bay of Bengal and over the Caribbean
- ΔD enhancements over Bay of Bengal and Caribbean seem not to travel further North in the central monsoon areas
- □ ΔD enhancements over Bay of Bengal and Caribbean seem to be unique features; no similar features exist in the SH during DJF => monsoon related
- □ Compare to convection areas over land:









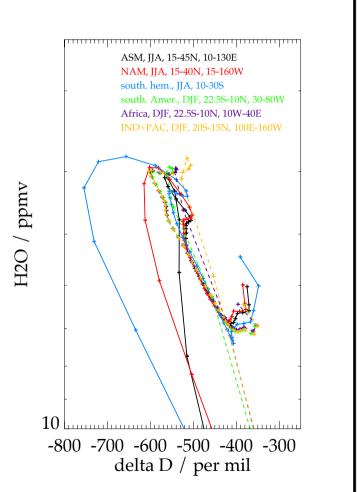
Preliminary conclusions and further work



- In contrast to analyses on pressure/altitude/T_{pot} levels the tropopause-following analysis demonstrates that water vapor (WV) minima are located below and at the monsoon tropopauses
- Above the tropopause WV is slightly (~ 0.5 ppmv) increased above the monsoon systems, compared to the rest of the tropics
- The Asian monsoon WV minimum is directly connected to that over the West Pacific/maritime continent
- AD helps to understand the phase transition processes during water WV uplift into the stratosphere
- The monsoon regions have about the same ΔD depletion as other convective regions over land, but are less depleted than the SH inner tropics, and H2O is higher
- Both findings hint towards uplift and deposition of ice in the TP region (± 2km) to play a more dominating role than in the SH inner tropics
- Convective processes over the Bay of Bengal and the Caribbean produce clearly enhanced ΔD values hinting towards strong ice lofting processes
- Future work: refine analysis using also radio occultation temperature data, studying the intra-seasonal variation, and resolving this mystery:

Preliminary conclusions and further work

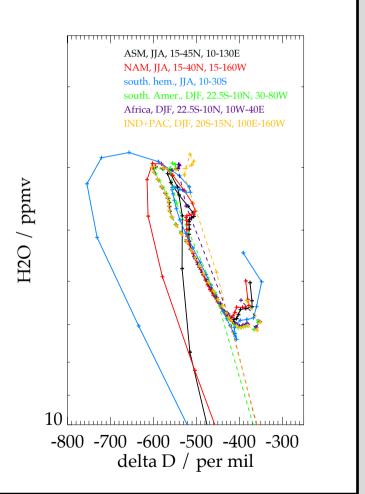
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Thank vou!



Gabriele Stiller. IMK-ASF

25 5 June 2017